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# Analysis of Groundwater Depletion and Recharge in Northwestern Bangladesh

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## ANALYSIS OF GROUNDWATER DEPLETION AND RECHARGE

#### IN NORTHWESTERN BANGLADESH

BY

MUHAMMAD S. KHAN

## A THESIS SUBMITTED IN PARTIAL FULFILLMENT OF THE REQUIREMENTS FOR THE DEGREE OF

## MASTER OF SCIENCE

IN

CIVIL AND ENVIRONMENTAL ENGINEERING

 $27104177$ UNIVERSITY OF RHODE ISLAND

 $\sim$ 

#### **ABSTRACT**

The decline of groundwater levels during the dry season was evaluated in a study area in northwestern Bangladesh. The feasibility of using recharge basins and recharge wells as a means for recharging the groundwater during the dry season was analyzed.

A two-dimensional finite-difference computer model of groundwater flow (MODFLOW) was used in conjunction with a field scale computer model of runoff from agricultural management systems (CREAMS) to evaluate the natural or artificial recharge to the groundwater from precipitation. Effects of artificial recharge from six recharge wells and four recharge basins were analyzed. Irrigation during the dry season utilizing the artificially recharged groundwater proved to be technically feasible.

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#### I. **INTRODUCTION**

The movement of groundwater is a part of the hydrologic cycle. An understanding of the role of groundwater in this cycle and the ability of man to manipulate it is mandatory if integrated analyses are to be performed to assess the problems associated with the watershed resources and enhance its availability. Groundwater should be treated as more than a human resource for it is also an important feature in the maintenance of ecologic balance. Its excess or deficiencies may lead to human and/or environmental problems, but at the same time, groundwater offers a medium for solutions to these problems.

Depletion of groundwater is a common phenomenon in the natural environment which may be the result of various artificial and natural circumstances, such as diversion of river flows or reduced recharge from precipitation. Depletion of groundwater may cause reduced growth of vegetation, posing adverse impact on the natural environment. Less availability of water for irrigation or drinking purposes and salt water intrusion along the coast line as a supplementary effect, may result from lowering of groundwater levels (Todd,1980; Freeze and Cherry,1979).

Attempts have been made to sustain the groundwater levels using various methods, such as recharging the groundwater artificially using recharge basins and wells (Kashef,1986), using subsurface dams (Hanson and Nilsson,1986; Suqio,Nakada, and Urish,1987) to maintain useable groundwater levels, and using irrigation return flows (Bouwer,1978). Use of recharge basins and recharge wells are widespread methods and their design, installation, operation, and maintenance do not require much effort. However, like any other artificial recharge method, implementation is not as simple as the theory holds. Major problems associated with the method include clogging of the recharge bed with finer particles (Kashef,1986) and air entrapment in the recharge wells (Freeze and Cherry,1979).

Evaluation of the potential impact of using artificial recharge methods to control groundwater levels is a complex task. An extensive analysis of the different processes involved in the hydrologic cycle is required to predict the effects of using such methods.

In order to analyze the feasibility of using recharge basins and recharge wells to recharge the groundwater artificially, numerical computer models may be used as tools to overcome the complexities involved in the analysis. However, a model has to be calibrated and validated with field

observations before it can be used to predict the effect of future modifications to the existing field conditions.

In addition to analyzing the feasibility of using an artificial recharge method, the computer models may also be used to assess various management options. For example, a model may be used to determine optimal locations for recharge, the most suitable engineering approach and management practice to augment the groundwater recharge.

The objectives of this study were 1) to evaluate the decline of groundwater levels in a study\_ area and 2) to analyze the feasibility of using recharge basins and recharge wells as a means for recharging the groundwater artificially.

Two computer models were involved in this study to evaluate the recharge to groundwater. A field scale model, CREAMS (Knisel,1980), was used to determine the deep percolation to groundwater. A finite-difference groundwater flow model, MODFLOW(McDonald and Harbaugh,1984), was used to determine the groundwater levels after recharge to groundwater takes place. Recharge basins and recharge wells were superimposed on the area to predict the possible increase in groundwater recharge.

#### II. **MODELINGAPPROACH**

### **2.1 Model Descriptions**

#### **CREAMS**

The hydrology component of the CREAMS model was utilized to predict the deep percolation to groundwater using daily precipitation records. This physically based model can be applied on a field scale. A field is defined (Knisel,1980) as a management unit having (1) a single land use, (2) relatively homogeneous soils, (3) spatially uniform rainfall, and (4) single management practices, such as conservation tillage or terraces.

The simulation of hydrologic response includes models for infiltration, soil water movement, and soil/plant evapotranspiration between storms. A time step of one day was used for evaporation and soil water movement between storms. The simulation for the period between storms provides prediction of amount of seepage below the root zone. A schematic representation of the processes involved in the model is shown in Figure 2. 1. A generalized flow chart of the simulation is presented in Figure 2.2.

Infiltration and runoff is predicted using SCS curve



Figure 2.1 : Schematic representation of CREAMS hydrolo option (Knisel, 1980).



Figure 2.2 : Generalized flow chart for CREAMS hydrology option (Knisel, 1980).

number technique (USDA,1972) from daily rainfall. An antecedent rainfall index is used to estimate the antecedent moisture as one of the three conditions (I-dry, II-normal, and **III-wet).** The relation between rainfall and runoff for these three conditions is expressed as a curve number (CN). Runoff is predicted using the SCS equation :

$$
Q = \frac{(P - 0.2s)^2}{P + 0.8s}
$$
 [2.1]

where  $Q$  is the daily runoff;  $P$  is the daily rainfall; and  $S$  is the retention parameter, all having dimensions of length. The retention parameter is related to soil water content with the equation:  $\mathbf{v}$ 

$$
s = s_{mx} \left( \frac{UL - SM}{UL} \right) \tag{2.2}
$$

where SM is the soil water content in the root zone, UL is the upper limit of soil water storage in the root zone, and  $s_{mx}$  is the maximum value of s. The maximum value of sis estimated with the I moisture condition CN using the SCS equation :

$$
S_{mx} = \frac{1000}{CN_{I}} - 10
$$
 [2.3]

where  $\texttt{CN}_\text{I}$  is the curve number (0 to 100) for moisture condit I. curve numbers for other moisture conditions and different management practices or hydrologic conditions have been updated based on experiments performed under different field conditions.

To account for the soil water distribution along the depth, the root zone is divided into seven layers and weighing factors (decreasing with depth).

Water that enters the soil, becomes either evapotranspiration, storage, or seepage below the root zone. The components of the water balance equation in the soil are evaluated with a time step of one day. The water balance can be expressed by the equation:

$$
SM_i = SM_{i-1} + F_i - ET_i - O_i + M_i
$$
 [2.4]

where  $F_i =$  infiltration from direct precipitation on day i  $ET<sub>i</sub>$  = plant and soil evapotranspiration on day i.  $O_i$  = seepage below the root zone on day i  $M_i$  = snow melt amount on day i  $SM = soil$  water storage in the root zone.

A snow accumulation and snow melt equation (Stewart et al.,1975) is used by the model to account for the snow melt component of the water balance equation.

The evapotranspiration (ET) component is computed by the . method followed by Ritchie (1972). Potential evaporation i computed by the equation :

$$
E_0 = \frac{1.28 \Delta H_0}{\Delta + \gamma} \tag{2.5}
$$

where  $E_0$  is the potential evaporation;  $\Delta$  is the slope of the saturation vapor pressure curve at the mean air temperature;  $H_0$  is the net solar radiation; and  $\gamma$  is a psychometric constant.  $\Delta$  is computed with the equation:

$$
\Delta = \frac{5304}{T^2} e^{(21.255 - 5304/T)}
$$
 [2.6]

where T is the daily temperature in degrees kelvin.  $H_0$  i calculated with the equation:

$$
H_0 = \frac{(1 - \lambda) (R)}{58.3}
$$
 [2.7]

where R is the daily solar radiation in langleys and  $\lambda$  is the albedo for solar radiation.

Potential daily soil evaporation is predicted with the equation:

$$
E_{SO} = E_0 e^{-0.4 (LAI)}
$$
 [2.8]

where  $E_{so}$  is the potential evaporation at the soil surface and LAI is the leaf area index defined as the area of the plant leaves relative to the soil surface area. Actual soil evaporation is computed in two stages. In the first stage, soil evaporation is limited only by the energy available at the surface, and thus is equal to the potential soil evaporation. Stage one upper limit of evaporation is computed with the equation :

 $\bullet$  .  $\bullet$ 

$$
U = 9 (\alpha_{s} - 3)^{0.42}
$$
 [2.9]

where U is the stage one upper limit in mm and  $\alpha$ , is soil evaporation parameter (ranges from 3.3 to 5.5 mm/ $\rm{d}^{1/2})$ . When the accumulated soil evaporation exceeds U, the stage two evaporative process begins. Stage two daily soil evaporation is predicted with the equation :

$$
E_g = \alpha_g [t^{1/2} - (t - 1)^{1/2}]
$$
 [2.10]

where  $\mathtt{E}_\mathtt{s}$  is the soil evaporation for day  $\mathtt{t}$ , and  $\mathtt{t}$  is th number of days since stage two evaporation began.

Plant evaporation (transpiration) is computed with the equations :

$$
E_p = \frac{(E_0) (LAI)}{3} , \quad 0 \le LAI \le 3
$$
 [2.11]

$$
E_p = E_0 - E_s , \quad LAI > 3
$$
 [2.12]

If soil moisture is limited, plant evaporation is reduced with the equation :

$$
E_{PL} = \frac{(E_p) (SM)}{0.25 \text{ FC}}, \quad SM \le 0.25 \text{ FC}
$$
 [2.13]

where  $E_p$  is the normal plant evaporation;  $E_{pL}$  is plant evaporation reduced by limited SM; **and FC is the field** capacity of the soil. Evapotranspiration, the sum of plant

and soil evaporation, can not exceed  $\texttt{E}_{\textup{0}}.$ 

Drought conditions are considered when the soil moisture falls below 15 bar amount or the permanent wilting point of the plant. Plant growth is stopped by holding the leaf area index constant until water becomes available.

Percolation or flow through the root zone is predicted using a soil storage routing technique (Williams and Hann,1978). The root zone is divided into seven layers or storages for routing. The routing equation is :

$$
O = \sigma (F + \frac{ST}{\Delta t}), \quad (F + \frac{ST}{\Delta t}) > FC
$$
 [2.14]

where Fis the infiltration or inflow rate; ST is the storage volume; *a* is the storage coefficient; and At is the routing interval (one day). If inflow plus storage does not exceed field capacity, FC, percolation is not predicted to occur. The storage coefficient is expressed by the equation:

$$
\sigma = \frac{2 \Delta t}{2t + \Delta t} \tag{2.15}
$$

where tis the travel time through a storage. Travel time is estimated with the equation

$$
t = \frac{SM - FC}{r_c}
$$
 [2.16]

where SM is soil water storage, and  $r_c$  is the saturated hydraulic conductivity of the soil.

Since each soil storage is subject to ET losses, the daily predicted ET must be distributed properly through the storages. A simulation of water use by root growth is expressed by the equation:

$$
u = u_0 e^{-4.16(RD)}
$$
 [2.17]

where u is the water use rate by the crop at root depth, RD, and  $u_0$  is the rate at the surface.

Extraction of water occurs from both surface and root zones in proportion to the relative root depth, which varies with leaf area index up to the maximum depth. Seepage from the root zone is predicted to occur when the moisture content exceeds the field capacity.

#### **MODFLOW**

In this study, MODFLOW(McDonald and Harbaugh,1984) was used to simulate the flow from external stresses, such as flow to and from wells, areal recharge, and flow through the bottom of the recharge basins. Groundwater flow within the aquifer is simulated using a block-centered finite-difference approach. Layers can be simulated as confined, unconfined, or a combination of confined and unconfined.

The three-dimensional movement of groundwater of constant density through porous earth material may be described by the

partial differential equation

$$
\frac{\delta}{\delta x} \left( K_{xx} \frac{\delta h}{\delta x} \right) + \frac{\delta}{\delta y} \left( K_{yy} \frac{\delta h}{\delta y} \right) + \frac{\delta}{\delta z} \left( K_{zz} \frac{\delta h}{\delta z} \right) - W = S_s \frac{\delta h}{\delta t} \quad [2.1]
$$

where x, y, and z are cartesian coordinates aligned along the major axes of hydraulic conductivity  $K_{xx}$ ,  $K_{yy}$ ,  $K_{zz}$ ; h is the potentiometric head (L); W is a volumetric flux per unit volume and represents sources and/or sinks of water  $(T^{-1})$ ; S. is the specific storage of the porous material  $(L^{-1})$ ; and t time (T) .

In general,  $S_{1}$ ,  $K_{xx}$ ,  $K_{yy}$ ,  $K_{zz}$  may be functions of space and h and W may be functions of space and time. Therefore, equation 2. 18 describes groundwater flow under non-equilibrium conditions in a heterogeneous and anisotropic medium.

The continuity equation is the basis for development of the groundwater flow equation in finite-difference form. The continuity equation can be stated as: the sum of all flows into and out of the cell must be equal to the rate of change in storage within the cell. Under the assumption that the density of groundwater is constant, the continuity equation expressing the balance of flow for a cell is:

$$
\sum Q_i = S_s \frac{\Delta h}{\Delta t} \Delta V
$$
 [2.19]

where  $Q_i$  is a flow rate into the cell  $(L^3T^{-1})$ ; S<sub>i</sub> is the specific storage defined as the ratio of volume of water which

can be injected per unit volume of aquifer material per unit change in head (L<sup>-1</sup>);  $\Delta V$  is the volume of the cell (L<sup>3</sup>); and  $\Delta h$  is the change in head over a time interval of length  $\Delta t$ . Thus a system of equations is developed to represent the flow system in each cell of the aquifer system.

A mathematical model of groundwater flow consists of equation 2.18 along with specification of flow and/or head conditions at the boundaries of an aquifer system and specification of initial head conditions.

In order to utilize the mathematical model, the aquifer system must be discretized into a finite number of cells. Figure 2.3 shows a spatial discretization of an aquifer system into a mesh of points termed nodes, forming rows, columns, and layers. Conceptually, nodes represent prisms of porous material, termed cells, within which the hydraulic properties are constant so that any value associated with a node applies to or is distributed over the extent of a cell. According to the block-centered formulation, the blocks formed by the sets of parallel lines are the cells; the nodes are at the center of the cells.

Different cell types are used to represent various types of boundaries. In general, the types of boundaries that may be imposed in the model include constant-head, no-flow,



Explanation

- Aquifer Boundary
	- Active Cell
	- 0 Inactive Cell
- $\Delta r_{\rm J}$ Dimension of Cell Along the Row Direction Subscripl (J) Indicates the Number of the Column
- $\Delta c_i$ Dimension of Cell Along the Column Direction. Subscript (I) Indicates the Number of the Row
- $\Delta V_{K}$ Dimension of the Cell Along the Vertical Direction. Subscript (K) Indicates the Number of **the Layer**

## Figure 2.3 : A discretized hypothetical aquifer syste (McDonald and Harbaugh, 1984).

constant-flow, and head dependent flow. An example of the use of no-flow and constant-head cells to simulate boundary conditions is shown in Figure 2.4. There are two types of boundaries that are integral to the model : an exterior noflow boundary at the edges of the model grid and internal boundaries consisting of no-flow and constant-head cells. Other boundary conditions such as specified flux can be simulated as a combination of no-flow boundaries and external stresses. However, it is not necessary to place no-flow boundaries at the exterior nodes of the grid.

The period of simulation is divided into a series of 'stress periods' within which all external stresses are constant. Each stress period, in turn, may be divided into a series of time steps. The system of finite-difference equations representing the aquifer system is formulated and solved to produce head at each node at the end of each time step. A generalized flow chart for the simulation is presented in Figure 2.5.

The computer program consists of a main program and a large number of highly independent subroutines called modules. These modules are, in turn, organized into 'packages' and 'procedures'. Table 2.1 shows the list of packages that constitute the model.



Explanation

Aquifer Boundary

Model Impermeable Boundary



~ Inactive Cell

Constant-Head Cell



**D Variable-Head** Cell

**Figure 2. 4** : **Discretized aquifer showing boundaries and constant head cells (McDonald and Harbaugh, 1984).**



DEFINE - Read data specifying number of rows. columns, layers, stress periods; and major program. op11ons

ALLOCATE - Allocate space in the computer to store dala

READ AND PREPARE - Read data which is constant throughout the simulation. Prepare the data by performing whate,e, caiculat,ons can be made **at** th,s s1age

STRESS - Determine the length of a stress period. and calculate terms to divide stress periods into time. steps

READ AND PREPARE - Read data which changes from one stress period to the next. Prepare the databy performing whatever calculations can be made at th,s stage

ADVANCE - Calculate length of time step and set heads at beginning of a new time step equal to heads. calculated for the end of the previous time step.

FORMULATE - Calculate the coefficients of the finite difference equations for each cell.

APROXIMATE - **Make** one cut **at** approx,mat,ng **a** solution to the system of finite difference equations.

OUTPUT CONTROL - Determine whether results should be written or saved on disk for this time step. Send signals to the BUDGET and OUTPUT procedures to indicate exactly what information should. be put out

 $BUDGET - Calculate terms for the overall volume$ met11c budget and calcu1a1e and **save** cell-by-cell flow terms for each component of flow.

OUTPUT - Print and save heads, drawdown and overall volumetric budgets in accordance with signals from OUTPUT CONTROL procedure.

**Figure 2.5 Generalized flow chart for MODFLOW (McDonald and Harbaugh, 1984).**

Packages are completely independent of each other. They can be added or removed without affecting other packages. There must, however, be a Basic package and a solver package.

Package Name	Abbreviation	Package Description
<b>Basic</b>	<b>BAS</b>	Manages the tasks that are part of the model as a whole.
Block- Centered Flow	BCF	Calculates terms of Finite- difference equations which represent flow.
Well	WEL	Adds terms representing flow to wells to the finite- difference equations.
Recharge	<b>RCH</b>	Adds terms representing areally distributed recharge to the finite-difference equations.
River	<b>RIV</b>	Adds terms representing flow to or from rivers to the finite-difference equations.
Drain	<b>DRN</b>	Adds terms representing flow to drains to the finite- difference equations.
Evapotrans- piration	<b>EVT</b>	Adds terms representing ET to the finite-difference equations.
General-Head <b>Boundaries</b>	<b>GHB</b>	Adds terms representing general- head boundaries to the finite- difference equations.
Strongly Implicit Procedure	SIP	Solver package for the system of finite-difference equations.
Slice- Successive Overrelaxation	<b>SOR</b>	Solver package for the system of finite-difference equations.

**Table** 2.1: List of **packages** of MODFLOW.

## **2.2 Input Data Requirements and sources**

In order to evaluate the existing hydrologic condition of the study area with the help of the computer models and to make any future predictions, information regarding different elements of the hydrologic cycle are required. Key data requirements for CREAMS and MODFLOWare listed in Table 2.2 and Table 2.3 respectively.

**Table 2.2 : Data requirements for CREAMS.**



Maps and soil profiles of the area were used to set up the study site. Topographic, groundwater contour, and soil association maps were used to identify the boundaries.

**Table 2.3: Data requirements for MODFLOW**

**Aquifer Properties: Hydrologic Data: Aquifer Stresses:** Areal dimension and boundaries Aquifer profile characteristics Storage coefficient Transmissivity Soil characteristics Stream discharge and water level Groundwater level Groundwater recharge Groundwater withdrawal

The primary sources of information regarding the study area are the reports on investigations conducted by different government and private organizations. Such organizations include Bangladesh Water Development Board (BWDB), Bangladesh Agricultural Research Council (BARC), Geological survey of Bangladesh (GSB), and Master Plan Organization (MPO). Raw data and information on detailed field investigations are available from the databases of some of these organizations.

MPO has been developing its own database collecting data from other sources. Most of the data used in this study was

available from this database. Agronomic and some climatologic data were available from BARC database. BWDB has detailed information about the monitoring of both the surface and groundwater. Well driller logs were available from GSB and MPO. Groundwater level monitoring data, Bore hole logs, and daily rainfall records at Shibganj (Fig. 3.1) are included in Appendices A, B, and C respectively.

In order to set up the models and to calibrate them, detailed information on soil, topography, and hydrology of the study area were required. In addition to the information directly related to the study area, general information regarding the study site were available from the instruction manuals of the models. Information from similar study sites was also considered. Using such information as a guide line, the models were more precisely set up and calibrated using data from field investigations. Reliability of the methods of collecting and recording of some data sources were sometimes questionable. Data from such sources were often crossexamined with a parallel source whenever necessary.

### **2.3 Modeling Procedure**

The primary intent of modeling the study area was to evaluate the recharge that is occurring to the groundwater.

Considering the complexities of the hydrologic processes both before and after recharge takes place, the modeling was carried out in two steps. In the first step, CREAMS was used to evaluate the percolation from the root zone to the groundwater. Then in the second step, MODFLOWwas used to determine the groundwater levels or heads.

A preliminary assessment of the problem was made without detailed information. This assessment defined the responses of the groundwater levels to the climate and to the boundary stream. Detailed raw data collected from the existing databases were consolidated to satisfy the requirements of the computer models. A base map was prepared to define the overall study area. Some other maps were associated to supplement the base map with information regarding different soil characteristics. The boundaries of the study area were selected considering different hydrologic information (such as groundwater divide, streams) and different soil and cropping classifications.

In order to evaluate the percolation from the root zone to the groundwater, it was necessary to account for different hydrologic processes that take place above, on, and below the ground surface. Part of the precipitation goes back to the atmosphere in the form of evaporation. The remainder is either carried out of the site as surface runoff, stored in

different forms, or infiltrated into the ground. Part of the infiltration is again evaporated as either soil or plant evaporation (transpiration). The magnitudes of all these elements depend largely on the soil, crop, and topographic characteristics. Two combinations of these characteristics were selected which would apparently yield extreme (minimum and maximum) percolation from the root zone. CREAMS was used to determine the percolation under these extreme conditions. The area was then divided into a number of different categories which would yield significantly different percolation.

A finite-difference grid of the modeled area was prepared to assign the percolation values from CREAMS and other relevant data to each node of MODFLOW.

MODFLOWwas calibrated and validated using two sets of different field data. During the calibration procedure, the initial values of transmissivity obtained from different sources were used as a guide line. These values were then adjusted to have a better agreement between the observed and modeled values of groundwater heads. After calibration, the model was validated with a different set of field data.

Four recharge basins and six recharge wells were selected to simulate artificial recharge to the groundwater.

Topographic and groundwater contour maps were utilized to select the most suitable locations of these basins and wells. MODFLOWwas used to determine the groundwater levels to quantify the magnitude of recharge.

Existing crop, crop calendar, and management practices of cultivation in the study area were modified to determine the impact of such modifications on the recharge to groundwater.

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#### III. **DESCRIPTION OF STUDY AREA**

## **3.1 Location**

The area selected for this study is located along the northwestern border of Bangladesh between latitudes 24°40'N &  $24°45'$ N and between longitudes 88°10'E & 88°16'E (Figure 3.1). The study area is situated in the upazilla (administrative unit comparable to a county in the U.S.) Shibganj of Nawabganj district. Specific features of the modeled site are discussed separately in chapter 4.

The Mohananda river runs along the eastern boundary of the study area. The Ganges river runs into Bangladesh from India approximately 40 kilometers {25 miles) away from the western boundary of the study area.

## **3.2 Climate**

Bangladesh has a tropical monsoon climate marked by sweltering temperatures and high humidity almost throughout the year. The country has four main seasons; Winter (December to February), Summer (March to May), Monsoon (June to September), and Autumn (October to November) (Mahmood, 1987).



Figure 3.1: Location of study area.

The study area is located along the northwestern border of the country, and is influenced by the Himalayan cold waves in Winter. However, in rare cases the temperature goes down to less than 41°F (5°C) and never touches the freezing point. Annual temperatures in this area range from 46°F (8°C) to 108°F (42°C) on the average. The coldest temperatures occur in the months of December and January and the warmest temperatures occur in April and **May.**

The average annual precipitation in this area is 59 inches (1500 mm). Rainfall occurs mostly during the southwest monsoon season from June to September. Tropical storms and thunderstorms are the sources of most of the precipitation in the monsoon season. Three years of monthly precipitation records measured at Shibganj are summarized in Table 3.1, the water year being April to March. The annual variation of rainfall is illustrated in Figures 4.3 and 4.4 and tabulated in Appendix c.

#### 3.3 **Land Use**

 $\bullet$ 

The study area is comprised of approximately 101.2 sq. mile (262.5 sq. km) land with 72% cultivated land. Of the total area, approximately 44% is highland, 24% is medium highland, 13% is medium lowland, 8% is very lowland, and 6% is
water bodies (MPO,1989). These lands are again divided into different land types depending on the flood phase (depth of flooding). Figure 3.2 shows the division of lands according to different flood phases.

	$1983 - 84$	$1984 - 85$	$1985 - 86$
April	20.6	0.0	27.2
May	58.4	83.7	89.0
June	81.3	242.6	226.0
July	243.2	336.0	349.4
August	192.3	368.9	189.2
September	162.2	352.5	330.2
October	195.6	226.3	110.5
<b>November</b>	0.0	0.0	0.0
December	26.5	0.5	0.0
January	27.7	0.0	0.0
February	17.8	5.1	0.0
March	0.0	6.4	0.0
<b>Total</b>	1023.6	1622.0	1321.5

**Table 3.1: Monthly precipitation (mm), Shibganj.**

Source: BWDB,1990

Generally, groups of small homesteads constitute the residential areas. Most of the roads are unpaved. Almost all the commercial and industrial activities take place at the upazilla headquarter, Shibganj. A land use map of the study area is shown in Figure 3.3. Table 3.2 explains the different land use associations.



Figure 3.2 : Different land types according to flood phase.



**Table 3.2 : Land use associations, Shibganj.**



### **3.4 Water Use**

There are three categories of water use in the study area; irrigation, domestic use, and industrial use. A major portion of the available water is used for irrigation. Most of the domestic usage is dependent on the available surface water from rivers, canals, and ponds. Drinking water is available from hand tube wells.

The water duty of the area for irrigation is 151 ha/Mm $^3\!$ . In other words, 151 ha land can be irrigated annually with 1  $m$ illion  $m<sup>3</sup>$  of water. The discharge per well is as follows: Shallow Tube Well (STW) -  $0.75$  to 1.0  $\mathrm{ft}^3\mathrm{/sec}$ ; Deep Tube Well (DTW) - 2.0  $\mathbf{ft}^3/\mathbf{sec.}$  The maximum pumping stress of thes wells are from April to May. Command areas of DTWs vary from 35 to 85 acres with an average of 54.71 acres; command areas

of STWs vary from 5 to 15 acres with an average of 9.81 acres **(MPO,** 1989) .

Rice and Rabi crops are the major crops produced in the study area. Some sugarcane, potato, and jute are also produced. There are three major categories of rice that are produced in this area : Aus, Amon, and Boro. Irrigation periods for these crops are as follows: Aus - mid March to June; Amon - July to October; HYV (High Yielding Variety) Boro  $-$  January to April; and Local Boro  $-$  December to mid April (BARC,1989).

# **3.5 Geology and Hydrogeology**

The study area is comprised entirely of one geomorphic unit - the flood plains being 97% of the total area. Active Gangetic flood plains and young meandering flood plains are the major physiographic units. The surface elevation from mean sea level ranges from 65 to 85 ft. in most parts of the area (MP0,1989).

The flood plains soils generally occupy a gentle landscape of low level to very gently sloping. The soil is mainly olive-brown, mixed grayish brown to olive brown, loamy to clay, silt loams or silty clay loams and are identified by

calcareous nature (MPO, 1989) . Figure 3. 4 shows different soil associations and land capability associations *in* the study area based on relief, age, and degree of weathering of surface sediments. Tables 3. 3 and 3. 4 explains the different soil and land capability associations. Table 4.2 describes the composition of the soil associations.

The study area constitutes a portion of the Indian platform of the Bengal Geosyncline. The subsurface stratigraphy of the area *is* presented *in* Table 3.5 based on drilled hole data obtained from the Geological Survey of Bangladesh (GSB).

Almost the whole of the study area *is* part of the active young Gangetic and mixed Gangetic and Mohananda flood plains and *is* underlain by unconsolidated recent and subrecent sequence of sand, silt, and clay.

The thickness of the upper silt and clay layer *is* about 49 ft. in the northwest and eastern side of the Pagla river and below 16 ft. in the rest of the area.

Maximum depth to groundwater table from the land surface varies from 20 to 30 ft. in the major portion of the area and 30 to 38 ft. in some small strips. The minimum depth to the groundwater table varies from 2 to 5 ft. throughout the area.





**Table 3.4 : Land capability associations, Shibganj.**



The minimum groundwater elevation from mean sea level is around 47 ft. near the river Mohananda and increases northwestward up to 65 ft. Groundwater fluctuates through a zone of 25 ft. in the central part of the area and decreases



**Table 3.5** .• **Hydrostratigraphy of the study area.**

to 7 to 10 ft. in the rest of the area. Water table decline during the dry season over a period of five years (1984 to 1988) varied from 0.43 to 10.10 ft. (MPO,1989). The general trend of groundwater movement is south towards the Mohananda river.

The transmissivity of the aquifer materials ranges from

35000 to 47500  $\mathsf{ft}^2\mathsf{/day}$  and specific yield value varies from 10 **to 15% (BWDB,1990; MPO,1989).**

Figure 3.5 shows the locations of groundwater monitoring wells and other investigation wells in the study area. Figure 3.6 shows two cross-sections of the soil strata as shown on Figure 3.5.







#### **IV. ANALYSIS**

### **4.1 Bydrologic Budgets**

Preliminary analysis of the study area involved estimating the overall water budget of the aquifer in terms of water loss or gain within a specified period of time. As a first approximation for the overall study area, the Thornthwaite method (Dunne and Leopold, 1978) was used for 1986. The potential evapotranspiration was estimated to be 34.61 inches. The average annual rainfall was 56.45 inches  $(1434$  mm).

The actual amount of recharge to the groundwater is, however, only a fraction of the difference between the rainfall and evapotranspiration due to loss in surface runoff. Identifying the annual storms in several groups and using the U.S. Soil Conservation Service Curve number technique (Dunne and Leopold,1978), the annual runoff volume was estimated to be 12. 49 inch. The remaining 9. 35 inch is the available recharge to the groundwater.

Regional groundwater maps were constructed from representative dry season and wet season water table data (Figures 4.1 and 4.2). An average variation of 10 ft was



Figure 4.1 : Regional groundwater map; June 16, 1986.





observed between the minimum and maximum water levels. Weekly water levels in the observation wells for 1986 are shown in Figures 4.3 and 4.4.

The rates of decline of the water levels in the wells during the dry season reflect the approximate groundwater flow when transmissivity is considered. The well hydrographs in general show very little difference in water table elevations at the beginning and the end of the year suggesting insignificant change in annual storage. The hydrographs reach their peaks toward the end of the monsoon and have a constant decline during the dry season.

In order to obtain a more precise hydrologic budget, evapotranspiration, runoff, and deep percolation from the root zone were estimated with CREAMS using daily precipitation records and mean monthly temperatures. For a silty loam soil with irrigated rice in 1986, predicted runoff was 13.85 inch, evapotranspiration was 51.10 inch, change in soil moisture in the root zone was 0.09 inch, deep percolation was 7.89 inch, and the applied irrigation to the root zone was 16.31 inch. Irrigation in this case means the amount of water actually supplied to the root zone by ponding either the rain water or the pumped water to meet the water demand of the plant. The significant difference between the evaporation estimates of Thornthwaite method and CREAMS prediction is due to the fact



Figure 4.3 : Weekly water levels in observation wells.





Figure 4.4 : Weekly water levels in observation wells.

that there is no correction for different vegetation types in Thornthwaite method (Dunne and Leopold,1978), whereas CREAMS uses Leaf Area Index values to consider different growth stages of the plant.

After considering different soil types and land uses, areal recharge resulting from the deep percolation was determined for MODFLOW predictions. Detailed discussion of these estimates are done in section 4.4. Considering a stress period of one month, for example June of 1986, the volumetric budget for the modeled site was as follows. Inflow to the aquifer from areal recharge was  $6.23 \times 10^6$  ft<sup>3</sup>; outfl through the pumping wells was  $1.97$  X  $10^6$  ft<sup>3</sup>, to the stream (constant head boundary) was 3.16 X 10 $^6$  ft $^3$ ; and the change in storage was  $1.10$  X  $10^6$   $ft^3$ .

# **4.2 Hydrology of Boundary Stream**

The Mohananda river runs along the eastern border of the study area. The rating curve for the river at station 210 (Tentulia) is shown in Figure 4.5. The discharge hydrograph of the river is shown in Figure  $4.6$ . The discharge hydrograph reaches its peak before the end of the wet season indicated. This may occur because the contributing watershed of the river is much larger than the study area whereas the precipitation



Rating Curve - Mohananda





Figure 4.6: Discharge hydrograph for the Mohananda river at Tentulia, Shibganj (MPO, 1989). Precipitation at Shibganj.

records shown reflect the pattern only in the more localized study area.

In order to determine whether the boundary stream is hydrologically connected to the aquifer, the river stages at station 210 during the dry and wet seasons were compared with the corresponding groundwater levels in the observation wells near the stream. It was observed that the groundwater levels during both the dry and wet season were at higher elevations than the river stages. However, the lowest water levels in the two wells RAJ75 and RAJ135 on the east side of the river were approximately at the same elevations as the river stage. Considering the depth of the river, it can be deduced that the river is hydrologically connected to the aquifer in the study area throughout the year. Moreover, observing the gradients of the groundwater table during the dry and wet seasons, the river can be identified as a gaining (effluent) stream.

# **4.3 Sensitivity of the Models**

A brief sensitivity analysis of the models was performed. This was necessary to evaluate performed response before the modeled site could be divided into areas yielding significantly different percolation and before the models could be calibrated.

#### **CREAMS**

One of the important parameters controlling the predicted percolation values was the curve number used to calculate runoff. Although listed values of curve number suggested for different hydrologic condition and cultural practice (Knisel,1980) were followed, the curve number was later modified. Considering the fact that ponding is required for the cultivation of rice, which would mean lower runoff across the dikes, a lower curve number than the suggested value was selected. Figure 4.7 shows the variation of predicted annual percolation values with curve number.

Five different soil types were selected to predict percolation. For an annual precipitation (1979) of 56.63 inch, for example, percolation from a clay soil was 10.20 inch; and from a sand loam soil was 18.26 inch. For these soil types, the predicted percolation values were most sensitive to the saturated hydraulic conductivity of the soils. Figure 4.8 shows the variation of predicted annual percolation values with saturated hydraulic conductivity.

The main crops in the modeled site were B.Amon (rice) and Rabi (winter crop). The predicted percolation values did not change significantly for different variety of rice or winter crops. However, the difference in land use caused a significant difference in the predicted values of percolation.



Figure 4.7 : Variation of predicted percolation with curve number.





#### **MODFLOW**

During the calibration procedure of MODFLOW, it was observed that the predicted head values depend largely on the estimated transmissivity of the aquifer and the areal recharge to the aquifer. Although the change in the predicted heads were more sensitive to percent change in recharge than to percent change in transmissivity, areal recharge to the aquifer was kept the same as the predicted percolation values of CREAMS while transmissivity estimates were modified. Figure 4.9 shows the variation of predicted head values with assumed transmissivity.

The iterative procedure used in calculating heads in MODFLOW prediction yields an approximation to the solution of the system of finite-difference equations for each time step (McDonald and Harbaugh, 1984). The rounding off error or truncation error is also associated with this procedure. However, even if a formal solution of the differential equations could be obtained, it would normally be only an approximation to the actual conditions in the field, because the hydraulic conductivity is seldom known with accuracy and uncertainties with regard to hydrologic boundaries are generally present.



Figure 4.9 : Variation of predicted head with transmissivity.

## **4.4 Groundwater Recharge in Existing Conditions**

An area of approximately 46.74 km $^2$  (5.03 X 10 $^8$  ft $^2$ ) was selected to evaluate the hydrologic budgets. Regional groundwater maps for a wet season and a dry season (Figures 4.1 and 4.2) were examined and compared to select the boundaries for MODFLOW. The modeled site with different land use and soil associations are shown in Figure 4.10. Groundwater elevations in the modeled site for a wet season and a dry season are shown in Figures 4.11 and 4.12 respectively. Considering the soil associations and the sensitivity of predicted percolation values, five different soil types were selected for CREAMS; namely, Clay (C), Silty clay (SiC), Silty clay loam (SiCL), Silty loam (SiL), and Sandy loam (SL). Physical soil properties including porosity, field capacity, and wilting point corresponding to each of the soil types were estimated from the listed experimental values (Knisel,1980; MPO,1989). The selected crops for simulation were B.Amon and Rabi. The leaf area index (LAI) values of the crops were calculated from the crop coefficient  $(K_c)$  curves for the corresponding crops (Doorenbos and Pruitt,1977). Initial and final water contents of soil for irrigation were also estimated (Jensen,1980).

Table 4.1 shows the predicted percolation values for 1986 determined with CREAMS for each soil type.



Figure 4.10 : Modeled site with different land use and sof



Figure 4.11: Groundwater elevations as of June 16, 1986.

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Figure 4.12 : Groundwater elevations as of October 20, 1986.

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**Table 4 .1** : **Predicted percolation values (inch) for different**



**soil types, 1986.**

These predicted values were then modified for each soil association based on the percent of each soil type in different soil associations. Table 4.2 shows the weighted percolation values for each soil association for 1986.

The weighted percolation values from CREAMS were used as areal recharge to the aquifer. Boundaries were selected for MODFLOW after examining the regional groundwater maps for a wet season and a dry season. The Mohananda river was selected as a constant-head boundary. The modeled site was divided



**Table 4.2** . **Weighted percolation values (inch) for different** • **soil associations, 1986.**

into 15 rows and 32 columns. The grid spacing was reduced near the stream and the proposed sites of recharge basins and recharge wells.

A steady state simulation of a dry month (June,1986) was performed to calibrate the transmissivity values at different nodes of the modeled site. Transmissivity values were first estimated from the pump test results (MPO,1989) and later adjusted for calibration. The groundwater contour map after calibration for June, 1986 is shown in Figure 4.13.

The calibration was verified with a steady state simulation of a dry month (October,1986). The corresponding groundwater contour map is shown in Figure 4.14. Table 4.3 shows the calibration and verification data for four locations within the modeled site.

A transient simulation of five wet months (June to October,1986) was performed next, with one-month stress periods. The resulting groundwater contour map is shown in Figure 4.15.

Location	Calibration		$1-\alpha$	Validation		$1-\alpha$
	Observed	Predicted		Observed	Predicted	
A	55.0	54.5	.98	64.3	67.6	.85
в	50.0	50.0		56.7	59.7	
C	57.0	56.2		65.6	68.5	
$\tilde{h}$ D	51.0	51.9		57.8	63.0	

**Table 4.3 : MODFLOWCalibration and validation data. water Levels (ft)**

1-a: t-test confidence level. Calibration: June, 1986; Validation: October, 1986.



Figure 4.13 : Calibrated groundwater contour; June, 1986.

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LEGEND:

**A,B..** Observation points for comparison

A.B..

Observation points for comparison


Figure 4.15: Transient simulation from June to October, 1986. Figure 4.15: Transient simulation from June to October,1986.

The hydrologic budgets evaluated for 1986 after calibration are shown in Table 4.4.





#### **4.5 Groundwater Recharge with Recharge Wells and Basins**

In order to evaluate the feasibility of using recharge wells and basins to increase the groundwater levels during the dry season, median percolation values were used rather than percolation from an average rainfall year like 1986. The median percolation value of a certain soil type would be that percolation, less than which would occur in 50% of the years. CREAMS was used to predict the monthly percolation in each soil type for 10 years for this purpose. An example of graphically determining the median percolation value is shown in Figure 4.16. Detailed results are tabulated in Appendix D.

The median percolation values determined graphically were modified for each soil association as discussed in section **4.4. A** transient simulation of the dry period {October to June) was performed to predict the resulting groundwater heads after the simulation period. The predicted groundwater contour map is shown in Figure 4.17.

Six recharge wells and four recharge basins were selected to recharge the groundwater during the dry season (Figure 4.18). Topographic and land use maps were examined to select the suitable locations of the wells and basins. Also total amount of runoff that can be captured during the wet season was estimated for each location using CREAMS. The evaporative loss from storage was also considered for each stress period.

The River Package in MODFLOWwas included to simulate the effect of having recharge basins. Conductance of the bed of the basins was calculated from the soil properties and basin -dimensions. The elevations of water levels in each basin for each stress period were determined considering the approximate amount of water recharged to the groundwater in the previous stress period and the evaporative loss from the surface of







Figure 4.17 : Transient simulation from October to June.

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Figure  $4.18$  : Locations of proposed recharge wells and basin

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water in the basin.

The rates of infiltration through the recharge wells in each stress period were determined from the amount of water available during the wet season. A constant infiltration rate was assumed as long as water is available. Therefore, some of the wells were inoperative toward the end of the dry season due to lack of water to infiltrate.

Effects of recharge wells and recharge basins were simulated separately first (Figures 4.19 and 4.20). Then the combined effects of recharge wells and basins were simulated (Figure 4.21). A comparison of Figures 4.19 and 4.20 with Figure 4.17 would suggest an increase in the groundwater levels when recharge wells and basins are used. Moreover, for this modeled site, recharge basins would increase the groundwater levels more than the recharge wells.

A simulation with recharge wells and basins for the wet season (June .to October) resulted in similar increase in groundwater levels (Figure 4.22).

In order to examine the feasibility of irrigating the land in the dry season to utilize the increased groundwater levels, the existing pumping wells were used along with 5 additional shallow pumping wells. Results of the simulation



Figure 4.19 : Predicted groundwater levels with recharge wells only; Simulation from October to June.

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Figure 4.21 : Predicted groundwater levels with recharge wells and basins; Simulation from October to June

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Figure 4.22 : Predicted groundwater levels with rechar<br>wells and basins; Simulation from June to Octobe

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are shown in Figure 4. 23. Table 4.5 shows the hydrologic budgets for this simulated year. Table 4.6 shows the simulation data for two schemes using artificial recharge.

**Table 4.5** • Hydrologic budgets for simulated year with recharge wells and basins.

		Volume $(ft^3)$
Aquifer	Areal Recharge	3.43 X 10 <sup>8</sup>
	Artificial Recharge	$1.03 \times 10^7$
	Discharge Wells	$2.23 \times 10^7$
	Flow to stream	$3.31 \times 10^8$
	Change in storage	$2.68 \times 10^3$

Table **4.6:** Simulation **data** for artificial recharge.

Location	Water Levels (ft)				
	Scheme 1		<b>Scheme 2</b>		
	Original	Simulation	Original	Simulation	
	A	55.0	66.5	55.0	61.4
	$\mathbf{B}$	50.0	58.0	50.0	56.4
	C	57.0	67.2	57.0	62.2
٠	D	51.0	64.0	51.0	58.5

Scheme 1 : Artificial recharge in existing condition. Scheme 2 : Artificial recharge with proposed irrigation.





#### **4.6 Modification in Cropping and Management Practices**

In order to examine any possible increase in percolation during the wet season due to change in'cropping or management practices, hypothetical situations of the modeled site were simulated using CREAMS with different crops and management practices. Non-irrigated B.Aus (rice) increased the annual percolation by 2.3 to 4.6% for different soil types. Winter crops, however, did not make any difference.

Although dikes built around the paddy fields reduce the area of cultivable land, ponded rice still proved to be suitable for groundwater recharge. In fact, any management practice that would lower the curve number and hence lower the runoff is suitable for natural recharge.

#### **V. CONCLUSIONSAND RECOMMENDATIONS**

In this study area, most of the rainfall occurs during the monsoon (July to October). The rest of the year is virtually rainless. As a result, the natural recharge to the groundwater occurs only during the wet season. The semiimpervious top layer (mostly silty clay) of the ground surface does not allow much water to infiltrate and most of the runoff is not available for natural recharge. Also, during the wet season, the wet soil reduces the infiltration rate significantly (Dunne and Leopold,1978).

During the peak dry season, the water table drops 16 to 33 ft. below the ground level, below the pumping suction limit of most of the shallow wells (BWDB,1990), thereby significantly reducing the irrigation capability and hence crop production. Installing more deep wells to withdraw water during the dry season would increase the irrigation capability during the dry season. But the same irrigation capability could be achieved with shallow wells throughout the year if the groundwater is artificially recharged to rise within the suction limit of the shallow wells during the dry season.

In this region, the percent of area under production of different varieties of rice have decreased approximately 17 to

44% depending on the variety over a period of 11 years (1978 to 1988) (MPO,1989) . This decrease is especially more prominent during the period 1981 - 1982 for the B.Amon variety and other irrigated rice. The reduction in agricultural land use availability for this period due to increase in residential and other land use was approximately 10%.

A change in the general trend of the dry season groundwater levels during the period 1981 - 82 were observed in the observation wells, especially in those towards the south of the area (Figure 5.1). These changes are apparently caused by the upstream diversion of the river flows in the Ganges (Abbas,1982; Begum,1987) which occurred during this time. However, analyses based on a larger watershed than that of this study has to be performed to determine any connection between these facts.

In this study, the installation of recharge wells and basins to increase the groundwater levels during the dry season was examined and proved to be feasible. The effects of using six recharge wells and four recharge basins were computer simulated to predict the increase in groundwater elevations during the dry season. Significantly higher water elevations than the existing condition were observed at the end of the dry season. In general, the water table in the area of artificial recharge was raised by 8 to 11 ft. Another



**Figure 5.1: Dry season water levels (77-88); well RAJ112. Precipitation at Shibganj.**

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simulation with increased irrigation {mostly with shallow pumping wells) showed that dry season pumping in conjunction **with** artificial recharge is feasible to increase annual rice production.

The practicality of installation and maintenance of recharge wells and basins, however, should be analyzed further. The method of construction of recharge wells may be different depending on the hydrogeologic conditions {Task Group of Artificial Ground Water Recharge,1965). Clogging of screens is the most serious problem in recharge wells {Olsthoorn,1982). Thus, screen open area and screen length must be optimal. Screens should be twice as long as for a withdrawal well pumping the same volume of water {Driscoll,1986). Other common practical problems include air entrapment and effect of injection and shut down periods {Sternau,1967; Rahman et al.,1969; Todd,1980; Bouwer,1978).

Runoff during the wet season should be captured in storage, and allowed to recharge groundwater later in the year through the recharge wells. The recharge basins consist of excavated basins in the ground or are created by dikes or levees surrounding the natural ground surface {Kashef,1986). Todd {1980), Task Group of Artificial Ground Water Recharge {1965), Bianchi and Muckel {1970), and Bouwer {1978) discussed the layout of a basin, or a series of basins, and methods of

their construction and maintenance. The most common problem associated with the maintenance of these basins is clogging of the recharge bed with fine particles. This problem may be overcome with periodic scraping of the top layer. Another practical problem associated with the recharge basins or the storage basin for recharge wells is the evaporation loss. Coverage (such as polyethylene sheet) could be used to minimize this loss. Underground storage tanks for recharge wells may also be feasible.

The high silt content in the runoff may be reduced significantly using both structural and non-structural Best Management Practices (BMP) (Land Management Project, 1990). Improving quality and controlling the quantity of runoff to receiving surface water and groundwater is a common purpose among these primarily preventive practices. Structural BMPs include sediment basins, artificial wet lands, and extended detention wet and dry basins. Non-structural BMPs include land use and site planning techniques, protection of natural buffer areas, and fertilizer management.

The approach followed in this study for a limited region in Bangladesh to augment the groundwater storage utilizing natural water supply may be used in other regions with similar hydrogeologic conditions. However, much depends on the rainfall magnitude and pattern. Precipitation should be

abundant during the wet season to recharge the groundwater during the dry season. Finally, the precise locations for installing the proposed systems have to be determined and further analyzed with more site specific field observation and evaluation.

#### **APPENDIX A**

Groundwater Level Monitoring Data



# RAJ73 1978 Well depth: 38.29 m; Dia: 0.04 m

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# RAJ75 1986 Well depth: 40.80 m; Dia: 0.04 m

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RAJ76 1986 Well depth: 9.60 m; Dia: 1.35 m

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### RAJ78 1986 Well depth: 8.53 m; Dia: 2.69 m

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### RAJ108 1986 Well depth: 8.76 m; Dia: 1.58 m R.L.24.99 m



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## RAJ110 1986 Well depth: 38.81 m; Dia: 0.04 m



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# RAJ112 1986 Well depth: 31.90 m; Dia: 0.04 m

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# RAJ135 1986 Well depth



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## **APPENDIX B**

Bore Hole Logs





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## GEOLOGICAL DATA RECORD CARD

BARIND INTEGRATED AREA DEVELOPMENT PROJECT,

### RAJSHAHI.

TEST WELL NO. 69 JL NO. 56 PLOT NO. 48 MOUZA Kansal Sixarpro LATITUDE 24°44'00" N LONGITUDE 88'11'57"E UPA-ZILL Shibgoni DISTRICT NONODC3.22 PROJECT BIAD project Rajebabi ORGANIZATION BADC COMMENCED 16.3.89 DRILLING CONDUCTED BY Graund Water Circle./ DATE COMPLETED  $22.7.89$ DATA RECORDED IN THE FIELD BY \_\_ Geologics', M. Shahedul Alson DRILLING EQUIPMENT USED Rig of XI PURPOSE Ground Alster Investive

TOTAL DEPTH (Metre / !!) 92.05m/302f/GROUND LEVELIMetre/MPWD DATUM 60.25 STATIC WATER LEVEL BELOW GL.(Metre) At HRS DATED 1.75 m. of 0700 h25 on 29.28







GEOLOGICAL DATA RECORD CARD BARIND INTEGRATED AREA DEVELOPMENT PROJECT. RAJSHAHI. WOUZA *Pitholitala* TEST WELL NO.  $20$  U NO.  $\frac{771}{20}$  PLOT NO. LATITUDE 84°40'15" N LONGITUDE 88°12'07"E UPA-ZILL Shindani DISTRICT AMAMAGANI PROJECT BLAD project Raishopi ORGANIZATION BADC COMMENCED 14.11.88 DRILLING CONDUCTED BY Grampd alater Crole. / DATE COMPLETED AS. 11.88 DATA RECORDED IN THE FIELD BY Jak tedden Khon Drilling sylv. DRILLING EQUIPMENT USED Rigno X1 PURPOSE Graund Water Toveshoust TOTAL DEPTH (Metre/!!) \_22.0500/308 GROUND LEVELIMETTE/MPWD DATUM 20.02 STATIC WATER LEVEL BELOW GL.(Melre) AI HRS DATED 3.05 m. at 0700 brs on 26.11.28 LOG COMPLETE-DEPTH THICKNESS LITHOLOGY COLOR AGE **FORMATION**  $\mathcal{L} \mathcal{L} \mathcal{A} \mathcal{Y}$ , liftle  $\mathcal{S} \mathcal{I} \mathcal{I}$ . <u>Grey</u>  $0 - 5.18$ ÷  $\overline{a}$  $5.12 - 14.63$ Fine SAND, trace very fine sond  $\overline{\mathcal{R}_{\mathbf{e}}}$  $q$  mich ξ CLAY, trace silt  $14.63 - 1533$ み。  $\boldsymbol{z}$ 4 L Medium SAND, trace coarse  $1505 - 40.54$  $\overline{\mathcal{R}}$  c.  $\tilde{\mathcal{L}}$  $\boldsymbol{\omega}$ <u>sand, fine sand g mica</u> Ú C SILT, pace very fine sand &  $40.54.60.96$  $\lambda$  o  $\overline{a}$ Concretion trace mico  $\boldsymbol{\varphi}$ Eine SAND, hace very fine sond<br>Silt & mica.  $\blacktriangleleft$  $7.33$  $\lambda o$  $S127$ , trace clay.  $72.33 - 92.65$ Zo.  $-4/5$ <u>c:/s</u> <u>RECAIRDES</u> Du Birector <u> BADC, shibeenigane</u>  $640 - 7.$ BNDB <u>Navnboanj.</u> Dhexa

GEOLOGICAL DATA RECORD CARD BARIND INTEGRATED AREA DEVELOPMENT PROJECT. RAJSHAHI. LATITUDE 24%0'00" N LONGITUDE BB°15'50"E UPA-ZILL Shibo'sqi DISTRICT AICLAND COOL PROJECT BIAD project Rojstabi ORGANIZATION BADC COMMENCED 16.3.89 DRILLING CONDUCTED BY Ground Clater Grale: J DATE COMPLETED  $33.3.99$ DATA RECORDED IN THE FIELD BY <u>Ceoloocet</u>, M. Shichecled Alum DRILLING EQUIPMENT USED\_Rigno. XI\_\_\_PURPOSE Groundwitter Invertigate STATIC WATER LEVEL BELOW GL.(Metre) At HRS DATED 3.81 m. 010700hrs on 22.9.29 **COMPLETE LOC** AGE FORMATION DEPTH THICKNESS LITH OLOGY **COLOR** SILT, with little cloy.  $0 - 10.97$   $10.97$  $27.$  Bro  $\omega n$  $\mathbb{R}^{\mathbb{Z}^2}$ Fine to Medium SAND, trace mica  $10.97 - 14.02$  3.05  $6794$  $\mathcal{E}$  $14.02 - 23.47$   $9.45$ SILT and very fine SAND, trace mico <u>Brown</u>  $\bm{\downarrow}$  $\boldsymbol{\mathcal{S}}$ e n  $73.47 - 28.55$  4.88 Medium to coorse sano, pace mica  $\overline{20}$  $4$   $11$  u  $vi$ Fine to Medium SAND, trace Cearse  $2835 - 35.92$  7.62  $6704$  $\checkmark$ sand a mica U Q  $5127$  with little clay.  $3592.59.29$   $14.57$  $500$  $777$  $50.29 - 54.25 - 3.96$  $\overline{\mathcal{X}^{\sigma}}$ Clayer SILT 54.25-104.4 49.99  $\ddot{\phantom{a}}$ ಸಾ  $C/5$  $C/5$ AE (BIEDR) Dy. Derector <u>BADC Shidganjzene.</u> GWD.D. BNDB, Dhowe <u>Navotdooj :</u>



**APPENDIX C**

Daily Precipitation Records (mm)

at Shibganj

 $\mathcal{L}_{\text{in}}$ 





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## **APPENDIX D**

Predicted Percolation Values (inch) from CREAMS

 $\frac{1}{2} \frac{Z_{\rm{eff}}}{\hbar}$ 

#### **Expected Percolation**

 $\mathbf{v}$ 

Soil: Clay, Crop: B. Aman & Rabi



 $\sim$  1  $\times$ 



### Soil: Silly clay; Crop: B. Aman &.Rabi

 $\mathcal{L}(\mathcal{A})$ 



#### Soil: Silty clay loam; Crop: B. Aman & Rabi

 $\label{eq:2.1} \begin{array}{c} \mathcal{L}_{\text{eff}}(\mathcal{L}_{\text{eff}}) = 0 \end{array}$ 



#### Soll: SIil loam; Crop: **B.**Aman & R,bl

 $\mathcal{L}(\mathbf{r})$ 

		1979	1980	1981	1982	1983	1984	1985	1986	1987	1988	Average
o	January	0	0	0	0	u	$\bf{0}$	0		0	o	
	February	0	۵	0	$\theta$	0	0	0		0	n	0
	March	o	0	0	$\mathbf{0}$	$\mathbf 0$	$\bf{0}$	$\mathbf{0}$	$\mathbf 0$	$\mathbf 0$	$\theta$	$\mathbf{0}$
	April	$\mathbf{0}$	a	0	0		0			0.194	a	0.019
	May	2.571		0	1.087	$\Omega$	$\Omega$			1.747	0.299	0.57
	June	$\mathbf 0$	0.942	0.211	5.259	$\mathbf 0$	0.437	1.164	0.046	$\mathbf 0$	5.076	1.314
	July	7.54	2.169	0.915	5.882	0.512	3.54R	0.827	4,926	6.284	4.443	3.704
	August	4.832	2.127	3.709	3.293	2.323	3.407	1.711	0.905	10.452	4.587	3.735
	September	0.09	3.767	4.291	1.12	0.534	10.145	6,744	2.596	2.529	5.123	3.694
	October	3.224	1.291	0.249	0.457	4.049	3.914	1806	6.993	0012	3.402	2.54
	November	$\bf{0}$	0.223	0	$\mathbf{0}$	$\mathbf 0$	0	0.011		0	O	0.023
	December	$\mathbf{0}$	0.12	0	0	0	$\bf{0}$	0	o	$\mathbf 0$	0	0.012
	<b>Total</b>	18.257	10.64	9.375	17.098	7.418	21.452	12.263	15.466	21.218	22.929	15.611

Soil: Sand loam; Crop: 8. Amon &.Rabi

 $\mathcal{L}^{\mathcal{L}}(\mathbf{R})$  . The contract of

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