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ANALYZING URBAN FLOOD DISASTERS IN

EMERGING MEGACITIES USING EARTH

OBSERVATIONS

BY

FARAH NUSRAT

A THESIS SUBMITTED IN PARTIAL FULFILLMENT OF THE

REQUIREMENTS FOR THE DEGREE OF

MASTER OF SCIENCE

IN

CIVIL AND ENVIRONMENTAL ENGINEERING

UNIVERSITY OF RHODE ISLAND

2018

MASTER OF SCIENCE

OF

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UNIVERSITY OF RHODE ISLAND 2018

ABSTRACT

The world is going through rapid urbanization resulting in cities turning into megacities. This rapid change turns into unplanned development in order to adapt to the growing population while the importance of the sustainability of the natural environment is neglected during the whole process. Along with the effects of climate change, flood disasters are becoming more frequent in megacities resulting in huge financial burden. Two driving factors behind urban flood disasters, anthropogenic and natural, are considered here: (i) Land Use and Land Cover (LULC) change, and (ii) intensity and frequency of precipitation. We focus on four major cities from across the world that are prone to chronic urban flooding problems: Houston, United States, Mexico City, Mexico, Jakarta, Indonesia, and Dhaka, Bangladesh. The aim of this study is to identify the main drivers behind flood disasters to improve disaster management and urban planning in these megacities. Utilizing the vantage of and recent advances in Earth Observations (EO) images and data, we assess urbanization patterns and associated hydrological changes for these cities. We found that LULC change is a principal driving factor behind urban flooding in Houston, Mexico City, and Dhaka. For Jakarta, both factors are equally important for urban flooding.

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Thank you all!

Farah

PREFACE

This thesis is submitted in a Manuscript Format. The first chapter is an introduction. The second chapter titled "Analyzing Urban Flood Disasters in Emerging Megacities Using Earth Observations" is prepared for submitting in Geophysical Research Letters, an American Geophysical Union (AGU) Journal.

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CHAPTER 1

Economic betterment is the root cause of the global phenomenon of rural to urban migration (Cohen, 2003). This mass movement of populations has been a vital part of the urbanization process from ancient times and continues till now (Lall, Selod, & Shalizi, 2006). Accepting the challenge of accommodating the increased population as part of the global urbanization process – cities are turning to megacities. A strong association is seen between population growth and land cover change (Dewan & Yamaguchi, 2009) while this urban expansion and growth lacks proper planning in necessary infrastructural development (Akanda & Hossain, 2012).

Urbanization is causing drastic changes in city layout and water infrastructures that are less resilient to natural disasters such as urban flooding. Land use land cover change

(K. A. Aderogba, 2012; K. Aderogba, Oredipe, Oderinde, & Afelumo, 2012; Odunuga, 2008), population growth, topography of an area, alteration in precipitation pattern and intensity, inadequate urban planning, and arbitrary solid waste disposal (Adeloye & Rustum, 2011; Lamond, Bhattacharya, & Bloch, 2012) are some of the reasons behind urban flooding. Among those, we selected two major factors, one anthropogenic and one natural, behind chronic urban flooding in emerging megacities: (i) land use land cover change – anthropogenic factor and (ii) changes of intensity and frequency of precipitation – natural factor.

The objective of this MS thesis is to uncover and understand the relative importance between (a) land use land cover change and (b) intensity and frequency change of precipitation for urban flooding. We found that LULC change is a principal driving factor behind urban flooding in Houston, Mexico City, and Dhaka. For Jakarta, both factors are equally responsible for urban flooding.

CHAPTER 2

Analyzing Urban Flood Disasters in Emerging Megacities Using Earth

Observations

By

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Is prepared for submission in Geophysical Research Letters, an American Geophysical Union (AGU) Journal

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Abstract

The world is going through rapid urbanization resulting in cities turning into megacities. This rapid change turns into unplanned development in order to adapt to the growing population while the importance of the sustainability of the natural environment is neglected during the whole process. Along with the effects of climate change, flood disasters are becoming more frequent in megacities resulting in huge financial burden. Two driving factors behind urban flood disasters, anthropogenic and natural, are considered here: (i) Land Use and Land Cover (LULC) change, and (ii) intensity and frequency of precipitation. We focus on four major cities from across the world that are prone to chronic urban flooding problems: Houston Metropolitan Statistical Area, United States, Mexico City, Mexico, Jakarta, Indonesia, and Dhaka, Bangladesh. The aim of this study is to identify the main drivers behind flood disasters to improve disaster management and urban planning in these megacities. Utilizing the vantage of and recent advances in Earth Observations (EO) images and data, we assess urbanization patterns and associated hydrological changes for these cities. We found that LULC change is a principal driving factor behind urban flooding in Houston, Mexico City, and Dhaka. For Jakarta, both factors are equally important for urban flooding.

Keywords: urbanization, urban flooding, Land Use and Land Cover, flood disasters, Earth Observations

4

1 Introduction

1.1 Background

Urban floods are an increasingly frequent and damaging environmental disasters across the globe. Due to high population growth, rapid urbanization and unplanned expansion is continuing in many regions of the planet and resulting in changes in land cover and land use (Dewan & Yamaguchi, 2009). In many developing regions of the world, this urban expansion and development lacks proper urban and regional planning and has led to large concentrations of substandard housing settlements with inadequate water, sanitation, and drainage infrastructure (Akanda & Hossain, 2012). As a result, a large portion of the world's urban dwellers has become vulnerable to natural disasters, especially during floods.

Changes in land use in many of these emerging megacities have exasperated hydrological processes and resulting flood events. Hydrological modifications from increased urbanization impact infiltration and evaporation at both temporal and spatial scales (Ali, Khan, Aslam, & Khan, 2011). Hence, runoff generation and flow patterns are altered, resulting in changes in the recurrence and severity of flooding (Ali et al., 2011). An increase in the volume of rainwater runoff and a decrease in natural storm water retention areas are also the consequences of illegal encroachment and illegal dumping of bigger populations also decrease the drainage capacity of natural canals. Land subsidence is also increasing at alarming rates in many megacities due to the unplanned extraction of groundwater (The World Bank, 2011). The gravitational capacity of natural drainage channels is hampered by land subsidence, which has

added more vulnerability to flooding and increased the risk of coastal flooding as well. The disaster risks in coastal cities are thus much greater when above factors are combined with sea-level rise and intense rainfall (The World Bank, 2011).

Change of land use is not only a physical process of transforming one land use to another but also is linked to the alteration of the social, political, economic, and cultural orientation of any society (Pangaribowo, 2018). Conversion of agricultural land to non-agricultural land is increasing to meet the land and housing requirements of growing populations, which has an impact on economy, society, and environment as well. Socio-economic factors, i.e., higher land price near the urban areas, the opportunity of diverse livelihoods, and chances of high income in urban areas, subsequently, have an impact on the land use change processes (Larasati & Hariyanto, 2018).

In this study, Houston Metropolitan Statistical Area, Texas, Mexico City, Mexico, Jakarta, Indonesia and Dhaka, Bangladesh have been chosen as example cities of urbanization and associated hydrological and land cover changes. These four major cities are all prone to chronic urban flooding problems, but each is chosen from four difference economic groups: a developed (United States), upper middle income (Mexico), and lower middle-income (Indonesia), and a recently graduated lower middle income from a least developed economy (Bangladesh). These countries have been selected to effectively compare the evolution of these trends and correlate the changes in each city's individual development contexts.

1.2 Objectives

Among many responsible driving factors, an area's topography, changes in land use and land cover (K. A. Aderogba, 2012; K. Aderogba et al., 2012; Odunuga, 2008), changes in precipitation intensity and frequency, urbanization and population growth, defective urban planning, and arbitrary solid waste disposal (Adeloye & Rustum, 2011; Lamond et al., 2012) are some important factors behind urban flooding. Among these, we investigate two major driving factors, one natural and one anthropogenic, behind the evolution of urban flooding in this study: (i) land use and land cover change, and (ii) intensity and frequency of precipitation. The goal of the study is to assess the relative importance to anthropogenic (land use land cover change) and natural (precipitation frequency and intensity) to urban flooding vulnerability and determine the strength and role of these drivers in the context of the four growing regions. The study covers the time period from 1979 to 2017, with a two-decade period (1997-2017) of overlapping availability of ground and Earth Observations (EO) of precipitation and land use land cover data.

2 Materials & Method

2.1 Study Area

In the United States, flooding is regarded as the number one among all natural disasters in terms of frequency as 28 out of 60 natural disasters were flood related between year 1980 to 2004 (Fang, Safiolea, & Bedient, 2006). Houston, Texas, ranked as the fifth largest metro area in the U.S. with a population over 7 million, is chronically vulnerable to large flooding disasters. The city is flood prone due to its close proximity to the Gulf of Mexico, the strong nature of Gulf Coast rainfall, rapid

urban growth resulting in more paved areas and roadways, the presence of clayey soils reducing infiltration, and high runoff rates along with mild slopes (Fang et al., 2006). Large flooding events in the city of Houston have occurred in 1989, 1992, 1994, 1998, 2001, 2003 (Fang et al., 2006), 2015 (Bass, Juan, Gori, Fang, & Bedient, 2016), 2016 and 2017 resulting in billions of dollars in damage and restoration costs. Over US\$52 billion costs due to flood damages at counties along the Gulf of Mexico in between 2000-2005 including US\$19 billion from the National Flood Insurance Program (NFIP) (Brody, Peacock, & Gunn, 2012). With current growth trends, the metro region of Houston, along with critical energy infrastructure, is likely to be highly vulnerable to future natural disasters.

Mexico City is located on the basin of Mexico, a lake basin, which is around 2260 meter above MSL. The city is surrounded by large mountains on three sides (Ochoa, Quintanar, Raga, & Baumgardner, 2015). This area had a large number of lakes and wetlands until the 1500s, and were subsequently drained and filled after the Spanish Conquest. The land cover of the city was a combination of shrubs and deciduous vegetation along with willows and pines on mountains before urbanization took place (Torres-Vera, Prol-Ledesma, & García-López, 2009). The growth of the city over the last 50 years can be divided into two groups: planned urban area for the middle and upper class population and unplanned urban areas near the periphery of the city for the poor and immigrants (Torres-Vera et al., 2009). Urbanization has intervened mostly in central and northern parts of the city whereas southern part is a blend of conserved forests, agricultural lands, wetlands and grasslands (Zambrano, Pacheco-Muñoz, & Fernández, 2018). The total urbanized area of the city consists nearly 20 million

people (Quintana-Belmares et al., 2018). A long history of illegal settlement and lack of demarcated land use between the center and the suburbs of Mexico City (Platt, 2010) led to unplanned urban development. Unceasing urban expansion along with climate change intensify spatial and temporal extent of flooding (Eakin et al., 2017). It has flooding history in the year of 1976, 1979, 1982, 1987, 1989, 1990, 1992, 1994, 1998, 1999, 2000, 2010 (Tellman et al., 2018). Most precipitation is observed between May and September with a variation in the northern and southern parts of Mexico City. The average annual precipitation in the southern areas is 1,200 mm, which is 600 mm in the northern areas (Romero Lankao, 2010).

Jakarta, the capital of Indonesia is also highly vulnerable to flooding disasters. Since 1980, Jakarta has undergone a dramatic transformation due to massive landdevelopment projects (Padawangi & Douglass, 2015). In the last fifty years, Jakarta's population has increased from 2.7 million in 1960 to about 9 million in 2007 (Budiyono, Aerts, Brinkman, Marfai, & Ward, 2015). This drastic increase has resulted in rapid changes in land use (Verburg & Bouma, 1999). Urban areas have become denser and only one-third of the city's area remains green and unpaved (Padawangi & Douglass, 2015). Real estate developers have invested in large geographical areas to maximize profits, resulting in large-scale land development projects, shifting existing land surfaces to urban areas. Floods have become a common consequence of the significant increase in paved area. In Jakarta, devastating flooding disasters occurred in 1996, 2002, 2007 and 2013, which inundated about 40% of the city in 2007. Such massive development has also led to significant subsidence in the northern parts of the Jakarta metro area, where a number of neighborhoods often experience coastal flooding and the old port area had to be protected by a seawall. Flood risk has dramatically increased due to population growth and a subsidence rate of 10 cm/year in some areas (Brinkman & Hartman, 2008).

Dhaka, the capital city of Bangladesh is an example of unplanned urbanization. It is one of the most densely populated cities in the world, with the highest growth rate in unplanned settlements (Akanda & Hossain, 2012). The population of Dhaka has increased from 3.44 million in 1981 (Dewan & Yamaguchi, 2009) to about 18 million in 2017 (The World Bank, 2017). Most of this growth have been absolved in unplanned settlements, where a large number of people have moved in from rural areas as economic migrants, climate refugees, and victims of natural disasters. Dhaka has an annual mean rainfall of 1920 mm and heaviest rainfalls occur between June and August (Hossain, Fien, & Horne, 2018). The city was originally developed in floodfree high lands, but the recent occupation of low-lying riparian suburbs around the city has drastically increased the flood vulnerability of the people (Adikari, Osti, & Noro, 2010). Low lying lands, rivers, canal, and water bodies are increasingly being filled to construct new accommodations on lands that previously worked as natural drainage channels (Hassan & Southworth, 2017). Artificial drainage is also hampered due to poor design of drains and sewer networks, unplanned construction, and dumping of uncollected wastes on the roadside (Yasmin & Rahman, 2017). Wetlands also operate as a recharge source of groundwater storage and allow drainage of extra precipitation that may otherwise cause urban flooding (IRIN, 2012). Thus, their recession has made the city more vulnerable to larger flooding events (IRIN, 2012). In addition, there is an embankment surrounding the Dhaka city to protect from river flooding. During

monsoon, river water levels are often higher than the city's water level inside the embankment, which creates hindrance in drainage by gravity (Mark, Wennberg, Van Kalken, Rabbi, & Albinsson, 1998).

2.2 Data

2.2.1 Landsat Images

All Landsat images for Houston, Jakarta and Dhaka were collected from the United States Geological Survey (USGS) Earth Explorer (USGS, 2017). Three years with decadal frequency were selected for comparison: 1997, 2007, and 2017. All images have a spatial resolution of 30 m. Only day-time images with cloud cover less than 10% were used for this analysis to allow the best visibility of land use and land cover. Landsat 5 imagery was used for 1997 and 2007, while Landsat 8 (launched in 2013) imagery was used for 2017.

2.2.2 Precipitation Data

We collected precipitation data (TRMM_3B42_Daily) from Tropical Rainfall Measuring Mission (TRMM) using the GES-DISC (Goddard Earth Sciences Data and Information Services Center) (Goddard Earth Sciences Data and Information Services Center (GES DISC), n.d.) for the specific Area of Interest (AOI) for the year of 1998 to 2017. The spatial resolution is 0.25° x 0.25° and temporal resolution is 1 day. The precipitation time-series is aggregated to several temporal scales (weekly, monthly, seasonal, and annual) for anomalies and trend analyses. Mann Kendall trend analysis and Sens's Slope is determined to better understand the statistically significant change in the precipitation trend. Trend of monthly maximum rainfall, total monthly rainfall, number of rainy days per month, total annual rainfall, maximum Consecutive Wet Days (CWD) in a year (annually maximum number of consecutive days with precipitation ≥ 1 mm), R10 (number of days annually when precipitation ≥ 10 mm), R20 (number of days annually when precipitation ≥ 20 mm), monthly and seasonal rainfall for each area were calculated. For longer-term trend analysis, CPC Global Unified Gauge-Based Analysis of Daily Precipitation (1979-2017) data (NOAA/OAR/ESRL PSD, Boulder, Colorado, n.d.) were collected for Houston and Jakarta. Bangladesh Meteorological Department Data (1953-2017) were collected for Dhaka. In CPC Global Unified Gauge-Based Analysis of Daily Precipitation, long range precipitation data are not available for Mexico. So TRMM data were used for trend analysis and SPI calculation of Mexico City.

2.3 Method

(a) <u>Detecting changes in LULC using Landsat Images:</u>

For Landsat 5 images, bands 1 to 5 were stacked to a single layer; and for Landsat 8 images, another layer was created by stacking bands 2 to 6. As we compared images in this study, we stacked bands of Landsat images with similar wavelengths (µm). A subset was created with each stacked layer according to the Area of Interest (AOI) and then unsupervised image classification with 40 classes is done. With the help of ERDAS IMAGINE 2016 software, each class was geo-referenced with Google Earth image of that particular period and assigned to a specific land cover: urban areas, vegetative cover, waterbodies, barren land, sand filled areas, future housing projects, and forest land. This helps to visualize the changes in land use/land cover over time. In the post classification process, urban areas were used as masks to detect other land

uses that were transformed to urban land use and the area calculation was done in acres.

(b) Creating land use/land cover change index:

The land use/land cover change index was created on the basis of infiltration capacity in this study. The lands which have more infiltration or drainage capability and transferred to paved areas are categorized with the highest index. The determination of the land-use index is shown in Table 1.

Initial Land use	Transformed	Land Use/Land
	Land use	Cover Change
		Index
Waterbodies	Urban Area	7
Vegetative Cover	Urban Area	6
Forest Land	Urban Area	5
Barren Land	Urban Area	4
Future Housing Projects	Urban Area	3
Sand filled Area (compacted)	Urban Area	2
Urban Area	Urban Area (No	1
	change)	

Table 1: Land use/land cover change index on the basis of infiltration capacity

(c) Non-parametric Mann-Kendall (MK) Trend Test

To determine the monotonic increasing or decreasing trend of climatological variables, non-parametric Mann-Kendall (MK) test (Mann, 1945) & (Kendall, 1955) ((Yu, Zou,

& Whittemore, 1993); (Douglas, Vogel, & Kroll, 2000); (Singh, Kumar, Thomas, & Arora, 2008)) is highly used due to its accommodating ability of missing values (Gajbhiye, Meshram, Mirabbasi, & Sharma, 2016). In this trend test, the null hypothesis (H₀) is there is no monotonic trend in the precipitation over time and the alternative hypothesis (H_A) is there is a monotonic trend (increasing or decreasing) available in precipitation over time. In any rainfall trend analysis, outliers will be there due to extreme rainfall events. These outliers have less impact (Birsan, Molnar, Burlando, & Pfaundler, 2005) on the result of this MK test as its statistics is based on positive or negative sign rather than any value (Gajbhiye et al., 2016). Here, "modifiedmk" package of RStudio software is used to determine the Mann-Kendal Trend and Sen's slope. We assume that the rainfall time series is independent.

$$S = \sum_{i=1}^{n-1} \sum_{j=i+1}^{n} sign(x_j - x_i)$$
(1)

where x_i and x_j are sequential data for the *i*th and *j*th terms, sign is the signum function, and *n* is the sample size.

$$sign(x_i - x_j) = \begin{cases} +1, & \text{if } x_j - x_i > 1\\ 0, & \text{if } x_j - x_i = 0\\ -1, & \text{if } x_j - x_i < 1 \end{cases}$$
(2)

The statistic *S* is nearly Gaussian when n = 18 with the mean E(S) and variance Var(S) of the statistic *S* given by

$$E(S) = 0, Var(S) = \frac{n(n+1)(2n+5)}{18}$$
(3)

If there is tie in the dataset, then Var (S) has to be adjusted and becomes

$$Var(S) = \frac{1}{18} \{ n(n-1)(2n+5) - \sum_{p=1}^{q} t_p \left(t_p - 1 \right) \left(2t_p + 5 \right) \}$$
(4)

The variable q and t_p are the number of tied groups and number of data values in the *p*th group, respectively. The standardized statistic (*Z*) for one-tailed test of the statistic *S* is given as follows:

$$Z_{mk} = \begin{cases} \frac{S-1}{\sqrt{Var(S)}} & if \ S > 0\\ 0 & if \ S = 0\\ \frac{S+1}{\sqrt{Var(S)}} & if \ S < 0 \end{cases}$$
(5)

An increasing trend is identified with a positive Z_{mk} and a decreasing trend is identified with a negative Z_{mk} . For 95% confidence interval (or significance level, α = 0.05, the critical Z value = \pm 1.96 (for a two tailed test) and for 90% confidence interval, |Z| = 1.65 (Q. Zhang, Xu, & Zhang, 2009). If |Z| > 1.96 (Q. Zhang et al., 2009), the null hypothesis can be rejected. For 99% confidence interval, |Z| = 2.58(Q. Zhang et al., 2009).

Sen's Slope

The magnitude of the trend change can be identified by a slope estimator β , which was first proposed by Sen (Sen, 1968) and then extended by Hirsch (Hirsch, Slack, & Smith, 1982). β is the median of overall all possible combinations of pairs for the whole dataset. The magnitude of trend was calculated predicted by the Sen's slope estimator with the slope T_i of all data pairs was computed as follows:

$$T_i = \frac{x_j - x_i}{j - i} \tag{6}$$

Where x_j and x_i are considered as data values at time j and I (j > i) correspondingly. The median of these N values of T_i is represented as Sen's estimator of slope. Sen's estimator is computed as

$$Q_{med} = \frac{T(N+1)}{2} \tag{7}$$

when N is odd, and it is considered as

$$Q_{med} = \frac{\left[T\left(\frac{N}{2}\right) + \frac{T(N+2)}{2}\right]}{2}$$
(8)

when N is even. At the end, Q_{med} is computed by two-sided test at 100 $(1 - \alpha)$ % confidence interval, and then a true slope can be obtained by the non-parametric test. A positive value of Q_i indicates an upward or increasing trend, and a negative value gives a downward or decreasing trend in the time series.

(d) <u>Standardized Precipitation Index (SPI)</u>:

Standardized Precipitation Index (SPI) was calculated for determining flood risks. SPI is generally used for monitoring drought (McKee, 1995) (McKee, Doesken, & Kleist, 1993) but has also been used to identify flood conditions where SPI can detect the development of soil-saturation conditions (Seiler, Hayes, & Bressan, 2002). SPI can be calculated for various temporal scales such as 1 month, 3 months, 6 months, 12 months, 24 months. SPI for 1 month is used here as the most relevant measure for appropriate soil moisture conditions for flooding. Long term precipitation data are required to calculate SPI, then the probability distribution function (Hayes, Svoboda, Wilhite, & Vanyarkho, 1999) is obtained from the data. Then, the cumulative distribution is transferred to normal distribution with a standard deviation of one, keeping zero as a mean value. Any positive SPI value indicates greater than median precipitation and vice versa. In 2000, Hayes et al. reported the interpretation of the Standardized Precipitation Index values (Hayes, 2000) into soil wetness measurements. The theoretical probability (Bonaccorso, Cancelliere, & Rossi, 2015) of occurrence of each interpretation derived from normal probability density function (Guhathakurta, Menon, Inkane, Krishnan, & Sable, 2018) is also given below:

SPI	Interpretation	Theoretical Probability
2.0 +	Extremely Wet	2.3
1.5 to 1.99	Very Wet	4.4
1.0 to 1.49	Moderately Wet	9.2
-0.99 to 0.99	Near Normal	68.2
-1.0 to -1.49	Moderately Dry	9.2
-1.5 to -1.99	Severely Dry	4.4
-2.0 and less	Extremely Dry	2.3

Table 2: Standardized Precipitation Index (SPI) ranges

(e) <u>Risk Index:</u>

Using the land use/land cover change index and the SPI, the risk index was created and transferred to a flood risk map for four cities for different months. Each SPI range has a probability of occurrence. Multiplying the SPI with LULC change index is giving us a probability which can be termed as risk index, which will give us an idea about how much urban flood risk any area has in terms of LULC change and precipitation alteration. The equation of risk index and the interpretation of the risk index range is given below.

$$Risk Index = Land Use/Land Cover Change Index X SPI$$
(9)

Table 3: I	Risk Inde	x and inter	pretation
------------	-----------	-------------	-----------

Risk Index Range	Interpretation
≤0	No Risk
0.01-6.00	Low Risk
6.01-12.00	Moderate Risk
12.01-18.00	High Risk
18.01-21.00	Extreme Risk

The flowchart of the methodology of this study is outlined below in Figure 1.



Figure 1: Land Use Land Cover Change Index using Landsat imagery (left column) and precipitation index using TRMM, CPC & BMD data (right column) were independently processed then findings are merged to identify the Urban Flood Risk.

3 Results

Land Use and Land Cover Data Analysis

In Houston, forest land and urban areas are the most changed land use and land cover categories from 1997 to 2017 (figure 2A-2C). According to figure 2A-2C, 33% of the total area was forest land in 1997 which was reduced to 20% in 2017. Vegetative cover was 29% of the total area in 1997 and has decreased to 24% in 2017. The percentage of barren land in 1997 was 3% of the total area, which has increased to 8% in 2017. Urban area, the most dominating land use in Houston, has increased from 31% to 45% between 1997 and 2017. Between 1997 to 2007, around 141,680 acres of different land covers were converted to urban areas, the changed land cover amount is 120,390 acres between 2007 and 2017.

In Mexico City, increase of barren land is observed from 1997 to 2007, which shows a decrease in 2017 (Figure 2D-2F). Urban area has increased from 33% to 37% of the total area from 1997 to 2017. Vegetative cover, which was 25% in 1997, is decreased to 22% in 2017. Around 35,000 acres and 54,000 acres of different land covers were converted to urban area in 1997 to 2007 and 2007 to 2017 respectively.

	Houston	n Metropolitan	Mex	ico City (a	res)	
		Area (acres)			ieo eny (a	
Land Cover/	year			year	year	year
Land use	1997	year 2007	year 2017	1997	2007	2017
Barren Land	29,475	53,843	71,783	51,235	60,159	44,274
Forest Land	299,820	267,997	181,988	201,980	177,932	195,658

 Table 4: Land use/ Land cover changes from 1997 to 2017 in Houston, and

 Mexico City

	Houston Metropolitan Statistical			Moy	ico City (a	oros)
	Area (acres)			IVIEX	neo eny (a	cres)
Land Cover/	year			year	year	year
Land use	1997	year 2007	year 2017	1997	2007	2017
Urban Area	282,092	348,966	405,239	206,679	208,183	232,631
Vegetative						
Cover	262,610	201,341	214,019	158,016	168,633	139,357
Water	10,584	15,286	19,252	3,814	6,816	9,804
Sand Filling	15,253	12,402	7,555			
Future						
Housing						
Projects						
Total Area	899,835	899,835	899,835	621,724	621,724	621,724

Extreme growth of urbanization is observed in Jakarta between the year 1997 to 2017. During this 20 years, urban areas have increased from 41% to 60% of the total area, which clearly reflects the haphazard urban expansion pattern. Apart from urban areas, vegetative cover and barren land are the two most changed land covers in this 20 years of span, both showing decreasing trends (Figure 2G-2I). 3% of the total area was designated as future housing projects in 2007 which is altered to urban areas in 2017. Around 48,000 acres of different land covers were transferred to urban areas between 1997 to 2007, which is increased to 69,000 acres between 2007 to 2017.

In Dhaka, vegetative cover and urban area are the two most changed land use and land covers among all the land use and land covers categories. From 2007, two new land uses and land covers are observed in Dhaka: sand filled areas and future housing projects, where both show increasing trends between 2007 and 2017. In 1997, vegetative cover consists 60% of the total area, which is reduced to 34% in 2017 (Figure 2J-2L). The percentage of urban area in 2017 has increased 1.5 times from that observed in 1997. The area that changed from different land covers to urban areas were 38,000 acres and 40,500 acres in 1997 to 2007 and 2007 to 2017 respectively. In 2017, 11% area of Dhaka city is designated as new housing projects apart from the existing urban area. These urban housing projects are not finished yet and will add more paved area in the city after completion.

Land Cover/	Jakarta (acres)			Jakarta (acres) Dhaka (acres)			
Land use	year 1997	year 2007	year 2017	year 1997	year 2007	year 2017	
Barren Land	39,544	18,134	3,586	7,442	11,908	12,239	
Forest Land	70,678	54,810	73,810	47,441	64,225	55,191	
Urban Area	145,052	159,137	210,043	52,845	65,959	80,532	
Vegetative							
Cover	64,588	81,860	35,691	189,984	147,469	106,299	
Water	29,857	26,800	26,590	17,220	16,282	17,792	
Sand Filling					3,565	7,524	
Future							
Housing		8,977			5,524.51	35,355.24	

Table 5: Land use/ Land cover changes from 1997 to 2017 in Jakarta, and Dhaka

Land Cover/	Jakarta (acres)				Dhaka (acres)	
Land use	year 1997	year 2007	year 2017	year 1997	year 2007	year 2017
Projects						
Total Area	349,719	349,719	349,719	314,933	314,933	314,933



Changes of Land Cover (Year1997-2007-2017)

Figure 2: Decadal Land use / Land cover changes from 1997 to 2017 in Houston (2A, 2B, 2C), Mexico City (2D, 2E, 2F), Jakarta (2G, 2H, 2I) and Dhaka (2J, 2K, 2L). The expansion of urban growth and transformation of other land uses and land covers to urban area in these cities with respect to times are shown in this figure.

Precipitation Analysis

(i) Detection of Mann-Kendall Trend Test and Sen's Slope

Mann-Kendall trend test was accomplished on the TRMM data, Bangladesh Meteorological Department (BMD) data, and Climate Prediction Center (CPC) Global Unified Gauge-Based Analysis of Daily Precipitation dataset for each of the study area. The test determines the trend change of monthly maximum rainfall, total monthly rainfall, number of rainy days per month, total annual rainfall, maximum Consecutive Wet Days (CWD) in a year (annually maximum number of consecutive days with precipitation ≥ 1 mm), R10 (number of days annually when precipitation ≥ 10 mm), R20 (number of days annually when precipitation ≥ 20 mm), monthly and seasonal rainfall for each area. For 95% confidence interval (or significance level, α = 0.05, the critical Z value = ± 1.96 (for a two tailed test) and for 90% confidence interval, |Z| = 1.65 (Q. Zhang et al., 2009). If |Z| > 1.96 (Q. Zhang et al., 2009), the null hypothesis can be rejected. For 99% confidence interval, |Z| = 2.58 (Q. Zhang et al., 2009).

In both locations of Houston (using CPC data), the monthly total precipitation in January, March, May, July, September, October, November and August (in location 2 only) is showing negative trend as both Z and Sen's slope is negative for these months but not statistically significant except the month of July for location 2. Same months showed negative trend in daily maximum rainfall analysis also, where July is showing statistically significant decreasing trend in both locations. In both locations, the trends of annual rainfall, R10, R20, total and maximum rainfall of summer, winter, spring, fall, dry and wet season are decreasing. Consecutive Wet Days (CWD) showed an

increasing trend in both locations, which are not statistically significant. For Mexico, maximum and total rainfall in dry months are showing positive trend and it shows decreasing trend for wet months. R10 and R20 is showing increasing trend, whereas CWD and annual rainfall in showing decreasing trend. None of them is statistically significant. The wet months of Mexico City is from May to October. Negative Z value and Sen's slope is observed in monthly total rainfall, daily maximum rainfall and number of wet days in those wet season months, which matches with the decreasing trend of annual rainfall and CWD. A long data series can help to better understand the trend. In Jakarta, the total and maximum rainfall in wet and dry period is showing positive trend along with positive Sen's slope. As one of the wet period months, January is only showing decreasing trend in monthly total, daily maximum and number of wet days on that month. February is showing statistically significant positive trend in monthly total and daily maximum. The trend of total and maximum rainfall for January-March is showing statistically significant positive trend. For Dhaka, the daily maximum rainfall trend for monsoon months (Jun-August) along with one pre-monsoon month (May) showing decreasing trend. Number of rainy days in monsoon months (June, August) showing decrease, which is statistically significant for June only. Maximum and total rainfall trend in monsoon and wet periods, annual rainfall and CWD are decreasing and only R10 and R20 is showing positive trend for Dhaka. The result of the Mann Kendal Trend Analysis and Sen's Slope are given below:
Trend	detection of st	ation data of	f Houston,	CPC Loc	ation 1 (197	9-2017)	
Mann-	-Kendall Trend	l Test & Sen	's Slope				
Indica	tors	Z-Value	Sen's	S	Var(S)	p	Tau
			slope			value	
	January	-1.016	-0.922	-85	6833.667	0.310	-0.115
	February	0.363	0.288	31	6833.667	0.717	0.042
all	March	-0.823	-0.991	-69	6833.667	0.411	-0.093
	April	0.169	0.132	15	6833.667	0.866	0.020
	May	-0.871	-0.556	-73	6833.667	0.384	-0.099
tal Rain	June	0.774	0.903	65	6833.667	0.439	0.088
ithly To	July	-1.669*	-2.404	-139	6833.667	0.095	-0.188
Mor	August	0.024	0.029	3	6833.667	0.981	0.004
	September	-0.387	-0.590	-33	6833.667	0.699	-0.045
	October	-0.629	-0.583	-53	6833.667	0.529	-0.072
	November	-0.556	-0.573	-47	6833.667	0.578	-0.063
	December	0.653	0.470	55	6833.667	0.514	0.074

Table 6: Mann Kendal Trend Analysis and Sen's Slope Analysis of Houston, CPC Location 1

Trend	detection of st	ation data of	Houston, G	CPC Loc	ation 1 (197	9-2017)	
Mann-	Kendall Trend	l Test & Sen	's Slope				
Indica	tors	Z-Value	Sen's	S	Var(S)	р	Tau
			slope			value	
	January	-0.073	-0.021	-7	6833.667	0.942	-0.009
	February	0.508	0.208	43	6833.667	0.611	0.058
	March	-0.097	-0.030	-9	6833.667	0.923	-0.012
	April	0.653	0.217	55	6833.667	0.514	0.074
	May	-0.944	-0.221	-79	6833.667	0.345	-0.107
Rainfall	June	0.895	0.337	75	6833.667	0.371	0.101
ximum l	July	-1.984*	-0.723	-165	6833.667	0.047	-0.223
aily Ma	August	-1.258	-0.477	-105	6833.667	0.208	-0.142
Ď	September	-0.532	-0.213	-45	6833.667	0.595	-0.061
	October	-0.460	-0.190	-39	6833.667	0.646	-0.053
	November	-0.435	-0.179	-37	6833.667	0.663	-0.050
	December	0.508	0.143	43	6833.667	0.611	0.058

Trend	Trend detection of station data of Houston, CPC Location 1 (1979-2017)								
Mann	-Kendall Trenc	l Test & Sen	's Slope						
Indica	tors	Z-Value	Sen's	S	Var(S)	р	Tau		
			slope			value			
	January	-2.327**	-0.118	-192	6739.333	0.020	-0.259		
	February	-0.292	0.000	-25	6740.333	0.770	-0.034		
	March	-1.458	-0.083	-121	6775.667	0.145	-0.163		
(1	April	-1.923*	-0.087	-158	6664.667	0.054	-0.213		
≥ 1mm	May	-1.253	-0.061	-104	6758.000	0.210	-0.140		
Rainfal	June	0.304	0.000	26	6774.667	0.761	0.035		
ıy days (July	-0.474	0.000	-40	6758.667	0.635	-0.054		
. of Rair	August	0.293	0.000	25	6732.333	0.770	0.034		
No	September	0.328	0.000	28	6762.000	0.743	0.038		
	October	-0.859	0.000	-71	6645.667	0.391	-0.096		
	November	-1.720*	-0.071	-142	6718.000	0.085	-0.192		
	December	-0.134	0.000	-12	6722.667	0.893	-0.016		
Annua	al Rainfall	-1.476	-6.175	-123	6833.667	0.140	-0.166		

Trend detection of sta	Trend detection of station data of Houston, CPC Location 1 (1979-2017)								
Mann-Kendall Trend	Test & Sen	's Slope							
Indicators	Z-Value	Sen's	S	Var(S)	р	Tau			
		slope			value				
Consecutive Wet	0.110	0.000	10	6676.667	0.912	0.013			
Days (CWD)									
R10	-1.079	-0.138	-90	6809.333	0.281	-0.121			
R20	-1.894*	-0.162	-157	6786.333	0.058	-0.212			
Total Rainfall of	-1.331	-2.707	-111	6833.667	0.183	-0.150			
Spring									
Total Rainfall of	-0.508	-1.536	-43	6833.667	0.611	-0.058			
Summer									
Total Rainfall of	-1.427	-2.437	-119	6833.667	0.153	-0.161			
Fall									
Total Rainfall of	-0.266	-0.299	-23	6833.667	0.790	-0.031			
Winter									
Maximum Rainfall	-1.282	-1.317	-107	6833.667	0.200	-0.144			
in Spring									

Trend detection of station data of Houston, CPC Location 1 (1979-2017)							
Mann-Kendall Trend	Test & Sen ³	's Slope					
Indicators	Z-Value	Sen's	S	Var(S)	р	Tau	
		slope			value		
Maximum Rainfall	-1.282	-2.005	-107	6833.667	0.200	-0.144	
in Summer							
Maximum Rainfall	-0.677	-0.694	-57	6833.667	0.498	-0.077	
in Fall							
Maximum Rainfall	-1.065	-0.781	-89	6833.667	0.287	-0.120	
in Winter							
Total Rainfall in	-1.766*	-6.282	-147	6833.667	0.077	-0.198	
Wet Season (MAR-							
NOV)							
Total Rainfall in	-0.266	-0.299	-23	6833.667	0.790	-0.031	
Dry Season (DEC-							
FEB)							
Maximum Rainfall	-1.355	-1.888	-113	6833.667	0.175	-0.152	
in Wet Season							
(MAR-NOV)							

Trend detection of station data of Houston, CPC Location 1 (1979-2017)						
Mann-Kendall Trend	Test & Sen ³	's Slope				
Indicators	Z-Value	Sen's	S	Var(S)	р	Tau
		slope			value	
Maximum Rainfall	-1.065	-0.781	-89	6833.667	0.287	-0.120
in Dry Season						
(DEC-FEB)						

95% confidence interval (**), 90% confidence interval (*)

Table 7: Mann Kendal Trend Analysis and Sen's Slope Analysis of Houston,

CPC Location 2

Trend	Trend detection of station data of Houston, CPC Location 2 (1979-2017)								
Mann-Kendall Trend Test & Sen's Slope									
IndicatorsZ-ValueSen'sVar(S)P-Tau									
slope S value									
fall	January	-0.847	-0.833	-71	6833.667	0.397	-0.096		
otal Rain	February	-0.097	-0.122	-9	6833.667	0.923	-0.012		
nthly Tc	March	-1.065	-1.087	-89	6833.667	0.287	-0.120		
Mo	April	0.290	0.340	25	6833.667	0.772	0.034		

Trenc	l detection of s	tation data of	f Houston,	CPC Loc	cation 2 (197	9-2017)	
Mann	n-Kendall Tren	d Test & Sen	's Slope				
Indica	ators	Z-Value	Sen's		Var(S)	P-	Tau
			slope	S		value	
	May	-0.774	-0.322	-65	6833.667	0.439	-0.088
	June	0.750	0.537	63	6833.667	0.453	0.085
	July	-2.226**	-2.472	-185	6833.667	0.026	-0.250
	August	-0.387	-0.378	-33	6833.667	0.699	-0.045
	September	-1.016	-0.978	-85	6833.667	0.310	-0.115
	October	-1.137	-0.938	-95	6833.667	0.255	-0.128
	November	-0.798	-0.604	-67	6833.667	0.425	-0.090
	December	0.992	0.713	83	6833.667	0.321	0.112
	January	0.000	-0.004	-1	6833.667	1.000	-0.001
Rainfall	February	0.460	0.117	39	6833.667	0.646	0.053
ximum]	March	-0.290	-0.152	-25	6833.667	0.772	-0.034
aily Ma	April	0.944	0.255	79	6833.667	0.345	0.107
D	May	-0.992	-0.193	-83	6833.667	0.321	-0.112

Trend	detection of s	tation data of	Houston,	CPC Loc	ation 2 (197	9-2017)	
Mann	-Kendall Tren	d Test & Sen	's Slope				
Indica	itors	Z-Value	Sen's		Var(S)	P-	Tau
			slope	S		value	
	June	0.968	0.313	81	6833.667	0.333	0.109
	July	-2.153**	-0.629	-179	6833.667	0.031	-0.242
	August	-1.718*	-0.393	-143	6833.667	0.086	-0.193
	September	-0.290	-0.105	-25	6833.667	0.772	-0.034
	October	-0.823	-0.356	-69	6833.667	0.411	-0.093
	November	-0.653	-0.221	-55	6833.667	0.514	-0.074
	December	0.895	0.252	75	6833.667	0.371	0.101
(1	January	-1.962**	-0.097	-162	6735.333	0.050	-0.219
l≥1mn	February	-1.134	-0.056	-94	6726.667	0.257	-0.127
(Rainfal	March	-1.630	-0.094	-135	6757.000	0.103	-0.182
ny days	April	-1.976**	-0.083	-163	6724.333	0.048	-0.220
). of Rai	May	-0.815	-0.037	-68	6753.333	0.415	-0.092
Nc	June	1.033	0.069	86	6765.333	0.301	0.116

Trend	l detection of s	tation data of	f Houston,	CPC Loc	cation 2 (197	9-2017)	
Mann	-Kendall Trend	d Test & Sen	's Slope				
Indica	ators	Z-Value	Sen's		Var(S)	P-	Tau
			slope	S		value	
	July	-0.243	0.000	-21	6751.000	0.808	-0.028
	August	0.000	0.000	1	6745.000	1.000	0.001
	September	-1.044	-0.069	-87	6781.667	0.296	-0.117
	October	-0.904	-0.032	-75	6703.000	0.366	-0.101
	November	-1.822*	-0.077	-150	6690.667	0.069	-0.202
	December	-0.757	0.000	-63	6708.333	0.449	-0.085
Annu	al Rainfall	-2.008**	-6.995	-167	6833.667	0.045	-0.225
Conse	ecutive Wet	0.061	0.000	6	6737.333	0.951	0.008
Days	(CWD)						
R10		-1.541	-0.185	-128	6791.333	0.123	-0.173
R20		-1.639	-0.143	-136	6780.667	0.101	-0.184
Total	Rainfall of	-1.331	-2.427	-111	6833.667	0.183	-0.150
Sprin	g						

Trend detection of station data of Houston, CPC Location 2 (1979-2017)								
Mann-Kendall Trend	Test & Sen	's Slope						
Indicators	Z-Value	Sen's		Var(S)	P-	Tau		
		slope	S		value			
Total Rainfall of	-0.847	-1.684	-71	6833.667	0.397	-0.096		
Summer								
Total Rainfall of	-1.790*	-3.125	-149	6833.667	0.073	-0.201		
Fall								
Total Rainfall of	-0.290	-0.242	-25	6833.667	0.772	-0.034		
Winter								
Maximum Rainfall	-1.185	-1.092	-99	6833.667	0.236	-0.134		
in Spring								
Maximum Rainfall	-2.056**	-2.621	-171	6833.667	0.040	-0.231		
in Summer								
Maximum Rainfall	-1.258	-1.228	-105	6833.667	0.208	-0.142		
in Fall								
Maximum Rainfall	-0.750	-0.459	-63	6833.667	0.453	-0.085		
in Winter								

Trend detection of st	ation data of	Houston, G	CPC Loc	ation 2 (1979	9-2017)	
Mann-Kendall Trend	Test & Sen	's Slope				
Indicators	Z-Value	Sen's		Var(S)	Р-	Tau
		slope	S		value	
Total Rainfall in	-2.032**	-6.693	-169	6833.667	0.042	-0.228
Wet Season (MAR-						
NOV)						
Total Rainfall in	-0.290	-0.242	-25	6833.667	0.772	-0.034
Dry Season (DEC-						
FEB)						
Maximum Rainfall	-1.935*	-2.480	-161	6833.667	0.053	-0.217
in Wet Season						
(MAR-NOV)						
Maximum Rainfall	-0.750	-0.459	-63	6833.667	0.453	-0.085
in Dry Season						
(DEC-FEB)						
	1			1		

95% confidence interval (**), 90% confidence interval (*)

Table 8: Mann Kendal Trend Analysis and Sen's Slope Analysis of MexicoCity, TRMM Location 5

Trend	Trend detection of Mexico City, TRMM (1998-2017)									
Mann-	Mann-Kendall Test & Sen's Slope									
Indicat	ors	Z-Value	Sen's	S	Var(S)	P-	Tau			
			slope			value				
	January	0.357	0.077	12	950.000	0.721	0.063			
	February	-0.519	-0.087	-17	949.000	0.603	-0.089			
	March	1.006	0.666	32	950.000	0.315	0.168			
	April	0.292	0.243	10	950.000	0.770	0.053			
Rainfall	May	1.460	2.633	46	950.000	0.144	0.242			
/ Total F	June	-0.422	-0.773	-14	950.000	0.673	-0.074			
Monthly	July	0.941	0.959	30	950.000	0.347	0.158			
	August	-0.357	-0.366	-12	950.000	0.721	-0.063			
	September	-0.876	-2.520	-28	950.000	0.381	-0.147			
	October	-1.914*	-2.354	-60	950.000	0.056	-0.316			
	November	-0.032	-0.015	-2	950.000	0.974	-0.011			

Trend of	Trend detection of Mexico City, TRMM (1998-2017)									
Mann-l	Kendall Test &	z Sen's Slope	2							
Indicate	ors	Z-Value	Sen's	S	Var(S)	P-	Tau			
			slope			value				
	December	-1.720(-0.210	-54	950.000	0.086	-0.284			
	January	0.097	0.017	4	950.000	0.922	0.021			
	February	-0.519	-0.054	-17	949.000	0.603	-0.089			
	March	1.395	0.300	44	950.000	0.163	0.232			
	April	1.655*	0.411	52	950.000	0.098	0.274			
fall	May	1.330	0.729	42	950.000	0.183	0.221			
ım Rain	June	-0.357	-0.102	-12	950.000	0.721	-0.063			
Maximu	July	-0.162	-0.051	-6	950.000	0.871	-0.032			
Daily	August	1.006	0.802	32	950.000	0.315	0.168			
	September	-1.460	-0.449	-46	950.000	0.144	-0.242			
	October	-0.357	-0.140	-12	950.000	0.721	-0.063			
	November	0.876	0.128	28	950.000	0.381	0.147			
	December	-0.487	-0.019	-16	950.000	0.626	-0.084			

Trend of	Trend detection of Mexico City, TRMM (1998-2017)									
Mann-	Mann-Kendall Test & Sen's Slope									
Indicat	ors	Z-Value	Sen's	S	Var(S)	P-	Tau			
			slope			value				
	January	0.132	0.000	5	915.000	0.895	0.026			
	February	-1.100	-0.056	-34	899.333	0.271	-0.179			
≥ 1mm)	March	0.624	0.068	20	926.000	0.532	0.105			
	April	-1.310	-0.154	-41	933.000	0.190	-0.216			
	May	1.113	0.222	35	933.000	0.266	0.184			
Rainfal	June	-0.491	-0.059	-16	932.667	0.623	-0.084			
ıy days (July	-0.426	0.000	-14	930.667	0.670	-0.074			
. of Raii	August	-2.819***	-0.317	-87	931.000	0.005	-0.458			
No	September	-0.982	-0.191	-31	933.667	0.326	-0.163			
	October	-2.280**	-0.500	-71	942.333	0.023	-0.374			
	November	-0.230	0.000	-8	930.000	0.818	-0.042			
	December	-0.036	0.000	-2	792.667	0.972	-0.011			
Annual		-0.681	-2.436	-22	950.000	0.496	-0.116			

Trend detection of Mexico City, TRMM (1998-2017)									
Mann-Kendall Test &	Mann-Kendall Test & Sen's Slope								
Indicators	Z-Value	Sen's	S	Var(S)	P-	Tau			
		slope			value				
Consecutive Wet	-0.361	0.000	-12	929.333	0.718	-0.063			
Days (CWD)									
R10	1.433	0.394	45	942.333	0.152	0.237			
R20	0.297	0.000	10	919.333	0.767	0.053			
Total Rainfall (NOV-JAN)	1.071	0.665	34	950.000	0.284	0.179			
Maximum Rainfall (NOV-JAN)	1.071	0.541	34	950.000	0.284	0.179			
Total Rainfall (FEB- APR)	1.200	1.401	38	950.000	0.230	0.200			
Maximum Rainfall (FEB-APR)	1.720*	1.003	54	950.000	0.086	0.284			
Total Rainfall (MAY-JUL)	0.746	1.642	24	950.000	0.456	0.126			

Trend detection of Mexico City, TRMM (1998-2017)									
Mann-Kendall Test &	Mann-Kendall Test & Sen's Slope								
Indicators	Z-Value	Sen's	S	Var(S)	P-	Tau			
		slope			value				
Maximum Rainfall	-0.552	-0.541	-18	950.000	0.581	-0.095			
(MAY-JUL)									
Total Rainfall	-1.330	-3.668	-42	950.000	0.183	-0.221			
(AUG-OCT)									
Maximum Rainfall	-0.941	-1.440	-30	950.000	0.347	-0.158			
(AUG-OCT)									
Total Rainfall in Wet	-0.876	-4.374	-28	950.000	0.381	-0.147			
Period (MAY-OCT)									
Total Rainfall in Dry	1.590	1.700	50	950.000	0.112	0.263			
period (NOV-APR)									
Maximum Rainfall	-1.071	-1.906	-34	950.000	0.284	-0.179			
in Wet Period									
(MAY-OCT)									
Maximum Rainfall	1.784*	0.943	56	950.000	0.074	0.295			
in Dry period (NOV-									

Trend detection of Mexico City, TRMM (1998-2017)								
Mann-Kendall Test & Sen's Slope								
Indicators	Z-Value	Sen's	S	Var(S)	Р-	Tau		
		slope			value			
APR)								

99% confidence interval (***), 95% confidence interval (**), 90% confidence

interval (*)

Table 9: Mann Kendal Trend Analysis and Sen's Slope Analysis of JakartaTrend detection of station data of Jakarta (1979-2017)

Mann	Mann-Kendall Test & Sen's Slope										
Indica	ators	Z-Value	Sen's	S	Var(S)	P-	Tau				
			slope			value					
	January	-0.121	-0.197	-11	6833.667	0.904	-0.015				
fall	February	2.976***	4.420	247	6833.667	0.003	0.333				
otal Rair	March	1.210	0.858	101	6833.667	0.226	0.136				
nthly Tc	April	0.750	0.638	63	6833.667	0.453	0.085				
Mo	May	0.677	0.570	57	6833.667	0.498	0.077				
	June	1.573	1.064	131	6833.667	0.116	0.177				

Trend	Trend detection of station data of Jakarta (1979-2017)									
Manr	Mann-Kendall Test & Sen's Slope									
Indic	ators	Z-Value	Sen's	S	Var(S)	P-	Tau			
			slope			value				
	July	0.798	0.878	67	6833.667	0.425	0.090			
	August	-0.992	-0.531	-83	6833.667	0.321	-0.112			
	September	-0.871	-0.644	-73	6833.667	0.384	-0.099			
	October	0.218	0.159	19	6833.667	0.828	0.026			
	November	1.185	1.038	99	6833.667	0.236	0.134			
	December	-1.331	-1.395	-111	6833.667	0.183	-0.150			
	January	1.573	0.423	131	6833.667	0.116	0.177			
	February	2.952***	0.909	245	6833.667	0.003	0.331			
Rainfall	March	1.355	0.227	113	6833.667	0.175	0.152			
ximum l	April	0.121	0.027	11	6833.667	0.904	0.015			
aily Ma	May	1.234	0.203	103	6833.667	0.217	0.139			
D	June	1.766*	0.314	147	6833.667	0.077	0.198			
	July	0.677	0.116	57	6833.667	0.498	0.077			

Trenc	Trend detection of station data of Jakarta (1979-2017)									
Mann	Mann-Kendall Test & Sen's Slope									
Indica	ators	Z-Value	Sen's	S	Var(S)	P-	Tau			
			slope			value				
	August	-0.750	-0.128	-63	6833.667	0.453	-0.085			
	September	-0.242	-0.082	-21	6833.667	0.809	-0.028			
	October	-0.073	-0.008	-7	6833.667	0.942	-0.009			
	November	0.024	0.011	3	6833.667	0.981	0.004			
	December	1.089	0.280	91	6833.667	0.276	0.123			
	January	-0.085	0.000	-8	6778.667	0.932	-0.011			
(1	February	1.515	0.074	125	6701.667	0.130	0.169			
l≥ 1mr	March	0.365	0.000	31	6758.333	0.715	0.042			
(Rainfal	April	1.034	0.054	86	6754.667	0.301	0.116			
ny days	May	0.707	0.029	59	6736.333	0.480	0.080			
. of Rai	June	0.194	0.000	17	6777.000	0.846	0.023			
No	July	0.461	0.048	39	6801.667	0.645	0.053			
	August	-0.668	-0.056	-56	6788.667	0.504	-0.076			

Trend	Trend detection of station data of Jakarta (1979-2017)									
Mann-Kendall Test & Sen's Slope										
Indic	ators	Z-Value	Sen's	S	Var(S)	P-	Tau			
			slope			value				
	September	-0.703	-0.067	-59	6798.333	0.482	-0.080			
	October	0.582	0.042	49	6794.333	0.560	0.066			
	November	0.451	0.000	38	6737.333	0.652	0.051			
	December	-0.558	-0.050	-47	6789.000	0.577	-0.063			
Annu	al Rainfall	1.573	12.553	131	6833.667	0.116	0.177			
Cons Days	ecutive Wet (CWD)	1.410	0.091	117	6769.000	0.159	0.158			
R10		0.788	0.200	66	6812.667	0.431	0.089			
R20		1.297	0.182	108	6801.333	0.194	0.146			
Total (JAN	Rainfall I-MAR)	2.468**	7.161	205	6833.667	0.014	0.277			
Total (APR	Rainfall R-JUN)	1.210	2.217	101	6833.667	0.226	0.136			

Trend detection of s	Trend detection of station data of Jakarta (1979-2017)								
Mann-Kendall Test & Sen's Slope									
Indicators	Z-Value	Sen's	S	Var(S)	P-	Tau			
		slope			value				
Total Rainfall	-0.290	-0.898	-25	6833.667	0.772	-0.034			
(JUL-SEP)									
Total Rainfall	0.435	1.277	37	6833.667	0.663	0.050			
(OCT-DEC)									
Maximum	1.984*	4.406	165	6833.667	0.047	0.223			
Rainfall (JAN-									
MAR)									
Maximum	0.653	0.359	55	6833.667	0.514	0.074			
Rainfall (APR-									
JUN)									
Maximum	-0.145	-0.182	-13	6833.667	0.885	-0.018			
Rainfall (JUL-									
SEP)									

Trend detection of s	Trend detection of station data of Jakarta (1979-2017)									
Mann-Kendall Test & Sen's Slope										
Indicators	Z-Value	Sen's	S	Var(S)	P-	Tau				
		slope			value					
Maximum	-0.097	-0.077	-9	6833.667	0.923	-0.012				
Rainfall (OCT-										
DEC)										
Total Rainfall in	2.516**	9.246	209	6833.667	0.012	0.282				
Wet Season										
(OCT-MAR)										
Total Rainfall in	0.508	1.952	43	6833.667	0.611	0.058				
Dry Season										
(APR-SEP)										
Maximum	1.645*	3.350	137	6833.667	0.100	0.185				
Rainfall in Wet										
Season (OCT-										
MAR)										
Maximum	1.161	0.839	97	6833.667	0.246	0.131				
Rainfall in Dry										
Season (APR-										

Trend detection of station data of Jakarta (1979-2017)								
Mann-Kendall Test & Sen's Slope								
Indicators	Z-Value	Sen's	S	Var(S)	P-	Tau		
		slope			value			
SEP)								

99% confidence interval (***), 95% confidence interval (**), 90% confidence

interval (*)

Table 10: Mann	NKendal Trend Analy	sis and Sen's Slo	pe Analysis of Dhaka
Trend dete	ection of BMD station	n data of Dhaka (1	1953-2017)

					Ň	,	
		Mann	ı-Kendall Te	est & Se	n's Slope		
	Indicators	dicators Z-Value		S	Var(S)	P-value	Tau
	January	-0.255	0.000	-42	25947.333	0.799	-0.021
ıfall	February	0.332	0.000	58	29532.000	0.740	0.029
otal Rain	March	1.179	0.233	204	29656.000	0.238	0.101
nthly Tc	April	0.348	0.240	61	29785.000	0.728	0.030
Moi	May	-0.336	-0.240	-59	29782.333	0.737	-0.029
	June	-0.829	-0.734	-144	29790.000	0.407	-0.071

Trend detection of BMD station data of Dhaka (1953-2017)										
	Mann-Kendall Test & Sen's Slope									
Ir	ndicators	Z-Value	Sen's slope	S	Var(S)	P-value	Tau			
	July	0.295	0.362	52	29790.000	0.768	0.026			
	August	0.220	0.140	39	29784.333	0.826	0.019			
	September	0.267	0.214	47	29789.000	0.790	0.023			
	October	0.238	0.171	42	29785.333	0.812	0.021			
	November	-0.843	0.000	-143	28349.667	0.399	-0.071			
	December	0.701	0.000	106	22418.667	0.483	0.053			
	January	-0.267	0.000	-44	25950.000	0.790	-0.022			
	February	-0.151	0.000	-27	29507.667	0.880	-0.013			
Rainfall	March	1.185	0.141	205	29633.667	0.236	0.102			
ximum]	April	0.157	0.009	28	29758.000	0.876	0.014			
aily Max	May	-0.151	-0.018	-27	29760.333	0.880	-0.013			
D	June	-0.852	-0.200	-148	29768.000	0.394	-0.073			
	July	-0.875	-0.213	-152	29770.667	0.381	-0.075			

	Trend detection of BMD station data of Dhaka (1953-2017)									
	Mann-Kendall Test & Sen's Slope									
In	ndicators	Z-Value	Sen's slope	S	Var(S)	P-value	Tau			
	August	-0.568	-0.147	-99	29781.000	0.570	-0.049			
	September	-0.435	-0.122	-76	29777.333	0.664	-0.038			
	October	0.713	0.141	124	29772.000	0.476	0.062			
	November	-0.778	0.000	-132	28346.667	0.437	-0.065			
	December	0.701	0.000	106	22419.333	0.483	0.053			
	January	-0.234	0.000	-37	23623.667	0.815	-0.018			
	February	0.708	0.000	120	28246.667	0.479	0.060			
l≥ 1mm	March	0.750	0.000	129	29127.000	0.453	0.064			
(Rainfal	April	0.782	0.000	135	29381.000	0.434	0.067			
ny days	May	1.027	0.000	177	29345.000	0.304	0.088			
of Rain	June	-2.978***	-0.063	-512	29452.000	0.003	-0.254			
No	July	0.397	0.000	69	29311.667	0.691	0.034			
	August	-1.152	0.000	-198	29240.667	0.249	-0.098			

	Trend detection of BMD station data of Dhaka (1953-2017)									
	Mann-Kendall Test & Sen's Slope									
Ir	ndicators	Z-Value	Sen's slope	S	Var(S)	P-value	Tau			
	September	1.055	0.020	182	29433.333	0.291	0.090			
	October	-0.625	0.000	-108	29294.667	0.532	-0.054			
	November	-0.678	0.000	-113	27254.333	0.498	-0.056			
	December	0.428	0.000	63	21019.667	0.669	0.031			
Ann	ual Rainfall	-0.145	-0.521	-26	29790.000	0.885	-0.013			
Cons Da	secutive Wet ys (CWD)	-1.489	-0.043	-257	29562.333	0.137	-0.127			
	R10	0.824	0.058	143	29709.667	0.410	0.071			
	R20	0.522	0.022	91	29670.333	0.601	0.045			
Tot (D	al Rainfall EC-FEB)	0.371	0.036	65	29729.667	0.711	0.032			
M Rair	laximum nfall (DEC- FEB)	0.522	0.057	91	29721.667	0.602	0.045			

Trend detection of BMD station data of Dhaka (1953-2017)									
	Mann-Kendall Test & Sen's Slope								
Indicators	Z-Value	Sen's slope	S	Var(S)	P-value	Tau			
Total Rainfall (MAR-MAY)	0.232	0.308	41	29787.000	0.817	0.020			
Maximum Rainfall (MAR- MAY)	-0.180	-0.106	-32	29783.333	0.857	-0.016			
Total Rainfall (JUN-SEP)	-0.214	-0.768	-38	29792.000	0.830	-0.019			
Maximum Rainfall (JUN- SEP)	-0.689	-0.875	-120	29788.000	0.491	-0.060			
Total Rainfall (OCT-NOV)	0.029	0.025	6	29790.000	0.977	0.003			
Maximum Rainfall (OCT-	0.023	0.000	5	29781.667	0.982	0.002			

Trend detection of BMD station data of Dhaka (1953-2017)								
Mann-Kendall Test & Sen's Slope								
Indicators	Z-Value	Sen's slope	S	Var(S)	P-value	Tau		
NOV)								
Total Rainfall in WET Season (MAY-OCT)	-0.504	-1.306	-88	29788.000	0.614	-0.044		
Maximum Rainfall in WET Season (MAY- OCT)	-0.904	-0.942	-157	29789.000	0.366	-0.078		
Total Rainfall in DRY Season(NOV- APR)	0.423	0.338	74	29786.000	0.672	0.037		

Trend detection of BMD station data of Dhaka (1953-2017)									
Mann-Kendall Test & Sen's Slope									
Indicators	Z-Value	Sen's slope	S	Var(S)	P-value	Tau			
Maximum									
Rainfall in DRY	0 267	0.162	17	29783 000	0 790	0.023			
Season(NOV-	0.207	0.105	Τ/	29783.000	0.790	0.025			
APR)									

99% confidence interval (***), 95% confidence interval (**), 90% confidence

interval (*)

(ii) SPI Calculation

For urban floods, we only consider SPI values for 1 month periods with values greater than 1, as -0.99<SPI<0.99 is near normal and negative values are typically considered for drought scenarios. Here, all the graphs are created for 1 month SPI to determine the soil moisture condition conducive to urban flooding. In the SPI graphs for Houston, it is clear that SPI values higher than 2, which indicates extremely wet conditions, are not frequent. Most of the SPI values are below 1 with some exceptional months. Year 1992, 1994, 1998, 2001, 2003, 2015 and 2017 are some of the flood years of Houston, which justifies their high SPI value. For Mexico City, only 4-5 months SPI values crossed 2. Moderately wet and very wet conditions are not recurrent. In Jakarta, SPI values greater than 1 are observed since 2007 (2007, 2010,

2012, 2013, 2017), which ranges from moderately wet to very wet. Extreme wet conditions are very few for Jakarta. For Dhaka, moderately wet to very wet condition is frequent since 1956. Extremely wet condition for Dhaka is observed very occasionally. The SPI graphs for all study areas are represented in Figure 3.



megacities

Validation of the LULC Analysis

The validation of the LULC analysis for Houston for year 2007 is done using 2006 National Land Cover Dataset (NLCD). The comparison of the total area and percentage is given below.

Houston 2007	Area	Percentage	Houston2006	Area	Percentage
	(acres)		(NLCD)	(acres)	
Parron Land	52.842	5.08	Barron Land	22264	2 / 1
Darren Lanu	33,043	3.90	Harbacous	32304	3.41
			Tierbaceous		
Forest Land	267,997	29.78	Deciduous Forest +	149491	15.76
			Evergreen Forest +		
			Mixed Forest +		
			Shrub/Scrub		
Urban Area	348,966	38.78	Developed Low	471023	49.64
			Intensity +		
			Developed Medium		
			Intensity +		
			Developed High		
			Intensity		
Vegetative	201,341	22.38	Hay/Pasture +	273435	28.82
Cover			Cultivated Crops +		
			Woody Wetlands +		
			Developed Open		
			Space		
Water	15,286	1.70	Open Water	14392	1.52
Sand Filling	12,402	1.38	Wetlands	8116	0.86
Future		_			_
Housing					
Projects					
Total	899,835	100.00	Total	948822	100.00

 Table 11: Validation of LULC Analysis for Houston

Urban Flood Risk Maps

The risk index calculated from equation (9) are subdivided into 5 classes. The risk index value ranges from -21 to 21. Any area is considered in no risk zone if the index value is negative or zero, as we are considering urban flood here. Other classes are: low risk (0.01-6), moderate risk (6.01-12), high risk (12.01-18) and extreme risk (18.01-21). Figure 4 represents the urban flood risk of study area for specific month and year.







(iv) Mexico City, August 2017

an



Figure 4: Urban Flood risk map of different wet season months of each study area to compare the change of risk between the year 2007 and year 2017

Validation of Urban Flood Risk Map with Federal Emergency Management Agency (FEMA) Flood Hazard layer

In our risk index calculation, we took LULC change and rainfall components only. Land elevation is another important factor that needs to be considered while calculating the urban flood risk index. For validating the urban flood risk map, we overlaid the map on the Federal Emergency Management Agency (FEMA) flood hazard layer for Houston. For better visualization, only a part of Houston is focused in Figure 5. FEMA updates their flood hazard data through Flood Insurance Rate Maps (FIRMs) (Xian, Lin, & Hatzikyriakou, 2015). FIRMs demarcate flood risk zones using hydrological and topographic survey outputs (Xian et al., 2015). The high risk and moderate risk areas of the urban flood risk maps match with 1% annual chance flood hazard, floodway, and 0.2% annual chance flood hazard area in most of the places with some exceptions. 1% annual chance flood hazard areas are defined as the areas that are situated on 100-year flood zones (Grineski, Collins, Chakraborty, & Montgomery, 2015). The FEMA flood map does not provide information on actual flood events but the probability of flooding (Grineski et al., 2015). A qualitative check on the two maps shows partial validation of our risk calculation approach, while there is room for improvement for areas that are prone to flooding due to distinct elevation changes such as rivers, streams, and bayous.



Validation of Urban Flood Risk Areas using FEMA National Flood Hazard Layer

Figure 5: Validation of Urban Flood risk map of Houston with FEMA flood hazard layer

4 Discussions and Conclusion

With the expansion of urbanization, vegetated soils convert to impervious surfaces that increase storm water flow and decrease both infiltration and natural storage (Wheater & Evans, 2009). Higher vegetative cover facilitates higher infiltration rate and quantity (Loch, 2000). The rate and magnitude of infiltration are dependent on the type, duration and intensity of precipitation, initial soil moisture content, soil type, evaporation, vegetation coverage, and terrain slope (G. Zhang, Qian, Wang, & Zhao, 2014). The soil composition of Houston is mainly the combination of fine sandy loam and clay, which has poor draining capacity (Muñoz, Olivera, Giglio, & Berke, 2018). Like Houston, the soil profile of Central Jakarta consists of alluvial clay in the form of soft to stiff (Hsiung, Yang, Aila, & Ge, 2018) and Dhaka city is a blend of Pleistocene clayey soils and Holocene clayey and sandy soils (Rahman, Kamal, & Siddiqua, 2018). As infiltration capacity of the clayey soil is less than that of sandy soil due to its smaller pore size, it is understood to be one of the main reasons that cause urban flooding in our study areas. In addition, rain on barren land compacts the upper layer of soil, creating hindrance in infiltration and causing excess runoff. Therefore, increasing amount of barren lands in these megacities are also responsible for urban flooding. According to Manning's equation, the velocity of the storm water flow is indirectly proportional with the roughness of the land surface (Leopold, Wolman, & Miller, 2012). Therefore, increasing paved smooth surfaces amplify the storm water flow more than any natural rough surface (Jacobson, 2011). Also, higher soil moisture has less ability to absorb extra runoff after precipitation. After analyzing the SPI graphs, it becomes clear that urban flooding is occurring despite of having low SPI
values, where low value indicates that the soil moisture is not high to be considered extremely wet. Rather, the values are indicating nearly normal to moderately wet soil moisture condition except some exceptional months.

Absence of zoning ordinance in Houston enables unplanned rapid increase of urban areas (Lynn, 2017). However, planners and developers have enough room to provide plans that can maximize urban and suburban vegetation within any development project (Conlon, Monaghan, Hayden, & Wilhelmi, 2016). The flat topography adds more difficulty in the flooding situation. Apart from planners, communities have started working on sustainable solution by changing unused golf courses to detention basins in the southeastern part of Houston to accommodate extra water after heavy precipitation (Landers, 2017). As reducing impervious layer is not easy inside cities, Low Impact Development (LID) practices can be helpful in reducing the excess runoff. These practices are used to manage storm water at the source by providing permeable pavements, bio-retention areas, and creating intermittent impervious surface (Damodaram et al., 2010). These could be potential remedies to decrease the heavy runoff due to impervious layers.

Due to the combined sewer system in Mexico City, volume of wastewater after heavy rainfall increases immensely. In base flow conditions, the waste water volume is 45 m³/s, which increases to 300 m³/s in peak flow conditions (Siemens, Huschek, Siebe, & Kaupenjohann, 2008). The pipe network of the combined sewer system is complex due to large difference in pipe diameter (0.30 m to 3.05 m). The system also generates sediment (Jiménez, Méndez, Barrios, Salgado, & Sheinbaum, 2004), which hampers the flow and creates more flood risk eventually for the city. The authority extracts 0.85 Mm³ of sediments (Jiménez et al., 2004) from the system every year and disposes as landfill, but cannot cope with the heavy rate of sedimentation in combined sewers.

Imprudent storm water drainage (Padawangi & Douglass, 2015) of Jakarta is another reason that is responsible for repetitive flood occurrence. Besides, insufficient finances to develop institutional capabilities, regulatory framework is also responsible for this situation (Kartez & Lindell, 1987). Changes of land ownership and extensive land development projects have influence on urban flooding (Walker, Whittle, Medd, & Walker, 2011). East flood canal project, the World Bank funded project named Jakarta Urgent Flood Mitigation Project / Jakarta Emergency Dredging Initiative (JUFMP/JEDI), proposed sea wall project is expected to be helpful in decreasing the flood risk in urban Jakarta (Padawangi & Douglass, 2015).

Apart from natural factors, stormwater management of these cities is also responsible for the situation. For example, Dhaka has only 30% and 38% coverage of sewerage and storm water systems, respectively (World Bank: BD: Dhaka Water Supply & Sanitation Project, 2017). Many areas of Dhaka have local combined sewer facilities, which cannot accommodate the excess runoff due to high-intensity precipitation. Surface runoff goes to underground sewer networks through catch pits. Inadequate intake capacity of catch pits or insufficient drainage capacity of sewer pipes cause surface flooding, which can contribute to urban flooding (Mark, Apirumanekul, Kamal, & Praydal, 2001).

The result of the study shows strong evidence that land use and land cover change (LULC) and insufficient water and drainage infrastructure development is mostly

accountable for urban flooding with moderate impact from precipitation alteration. Urban flood can occur any time but the frequency of occurrence is higher in wet periods of any area. The rainfall trends in wet periods of Houston, Mexico City and Dhaka are negative and for Jakarta, it is positive. It implies that land use land cover change is the main driving factor behind urban flooding in Houston, Mexico City, and Dhaka. For Jakarta, both factors are equally important for urban flooding.

Before approving any area as urban area, planners should test soil characteristics, which play a vital role in infiltrating floodwater and excess runoff. Accuracy metrics for LULC change analysis should be added in future analysis. In addition to protecting as much land as possible to preserve natural hydrological and drainage characteristics, installation of high capacity pumping stations, accommodating Low Impact Development (LID) practices should be incorporated at planning and implementation levels. Natural canal excavation to increase capacity, reclaiming illegally filled canals, separate sewage and storm water drainage system, and provision of retention basins and rainwater harvesting can further reduce the intensity of urban flooding conditions in developing cities. Strict law enforcement is also required in order to track and stop the illegal landfilling of the natural drainage system. Proper zoning is necessary to stop haphazard urbanization. As the world is rapidly urbanizing, steps to identify and reduce urban flooding disasters with the assistance of Earth Observations based analysis in the fastest growing megacities should be encouraged and adopted.

Appendix I: Map of Unsupervised Image Classification of the study areas

Land Cover Changes in Houston, USA (Year 1997-2007-2017)

Year 2007 Year 1997 15 30 60 0 Kilometers Legend E **Barren** Land Forest Land Sand filled areas **Urban Areas** Vegetative Cover Waterbodies

Year 2017

Figure: Land Cover changes in Houston, USA

Land Cover Changes in Mexico City (Year 1997-2007-2017)



Year 2017

Figure: Land Cover changes in Mexico City, Mexico

Land Cover Changes in Jakarta, Indonesia (Year 1997-2007-2017)



Figure: Land Cover changes in Jakarta, Indonesia

Land Cover Changes in Dhaka (1997-2007-2017)



Figure: Land Cover changes in Dhaka, Bangladesh

Appendix II: Post-classification land-cover change detection of the study areas

Land Use/Land Cover Change Index in Houston, Texas (Year 1997-2007 & 2007-2017)



Changes From Year 2007 to Year 2017

Land Use/Land Cover Change Index in Mexico City, Mexico (Year 1997-2007 & 2007-2017)



Changes From Year 2007 to Year 2017

Land Use/Land Cover Change Index in Jakarta, Indonesia (Year 1997-2007 & 2007-2017)



Changes From Year 2007 to Year 2017

Land Use/Land Cover Change Index in Dhaka (Year 1997-2007 & 2007-2017)



Changes From Year 2007 to Year 2017

Land Use/Land Cover changed	Houston				Mexico City			
	year 1997		year 2007		year 1997		year 2007	
	to 2007	Percentage	to 2017	Percentage	to 2007	Percentage	to 2017	Percentage
barren land to urban	8,707.88	6.15	19,342.10	16.07	19,440.23	55.74	12,157.67	22.38
forest land to urban	64,048.87	45.21	53 <i>,</i> 895.66	44.77	5,796.95	16.62	18,400.74	33.87
urban area (no change)	207,286.17	0.00	284,848.77	0.00	170,949.90	0.00	178,305.75	328.22
vegetative cover to urban	56,029.96	39.55	38,263.96	31.78	9,527.40	27.32	21,930.36	40.37
waterbodies to urban	530.63	0.37	1,474.04	1.22	112.53	0.32	1,836.98	3.38
sand filled area to urban	12,362.06	8.73	7,413.98	6.16	-	0.00	-	0.00
future housing projects to urban	-	0.00	-	0.00	-	0.00	-	0.00
total area (urban not included)	141,679.41		120,389.74		34,877.11		54,325.75	

Appendix III: Transformation of different land use/ land covers to urban areas

Land Lise (Land Cover shanged	Jakarta				Dhaka			
to Urbon Area	year 1997		year 2007		year 1997		year 2007	
to orbail Area	to 2007	Percentage	to 2017	Percentage	to 2007	Percentage	to 2017	Percentage
barren land to urban	15,176.23	31.51	10,529.74	15.32	1,348.60	3.54	3,357.27	8.28
forest land to urban	11,573.89	24.03	20,188.35	29.37	9,222.72	24.21	14,111.64	34.80
urban area (no change)	110,970.56	0.00	141,308.13	0.00	27,863.41	0.00	39 <i>,</i> 985.73	0.00
vegetative cover to urban	20,269.97	42.08	32,152.75	46.78	25,767.58	67.64	19,908.57	49.10
waterbodies to urban	1,146.45	2.38	2,547.09	3.71	1,756.47	4.61	1,608.58	3.97
sand filled area to urban	-	0.00	-	0.00	-	0.00	414.77	1.02
future housing projects to urban	-	0.00	3,317.02	4.83	-	0.00	1,146.00	2.83
total area (urban not included)	48,166.53		68,734.95		38,095.37		40,546.83	

Satellite	Sensor	Path	Acquisi		Wavelength	Useful for Mapping
		/Ro	tion	ion	(µm) of the	
		W	Date	olut m)	stacked	
				ses.	Spectral	
				н	Bands	
Landsat	Thema	137/	January	30	Blue (Band 1:	Bathymetric mapping,
5	tic	44	26,		0.45–0.52)	distinguishing soil from
(1984–	Mappe		1997			vegetation, and deciduous
2012)	r (TM)					from coniferous
						vegetation(Hugh-Jones,
						1989)
					Green (Band	Emphasizes peak
					2: 0.52–0.60)	vegetation, which is
						useful for assessing plant
						vigor
					Red (Band 3:	Discriminates vegetation
					0.63–0.69)	slopes
			January		Near-infrared	Emphasizes biomass
			22,		(Band 4:	content and shorelines
			2007		0.76–0.90)	
					Shortwave-	Discriminates moisture
					infrared 1	content of soil and
					(Band 5:	vegetation; penetrates thin
					1.55–1.75)	clouds
Landsat	OLI_T	137/	January	30	Blue (Band 2:	Bathymetric mapping,
8 OLI /	IRS	44	17,		0.452–0.512)	distinguishing soil from
TIRS			2017			vegetation, and deciduous
(2013–						from coniferous
Now)						vegetation
(Gorelic					Green (Band	Emphasizes peak
k et al.,					3: 0.533–	vegetation, which is
2017)					0.590)	useful for assessing plant
						vigor
					Red (Band 4:	Discriminates vegetation
					0.636-0.673)	slopes
					Near-infrared	Emphasizes biomass
					(Band 5:	content and shorelines
					0.851–0.879)	

Appendix IV: Specifications of Landsat TM and OLI images

Shortwave-	Discriminates moisture
infrared 1	content of soil and
(Band 6:	vegetation; penetrates thin
1.566–1.651)	clouds

(Pal & Ziaul, 2017)

Appendix V: Location of the TRMM stations and CPC gauges for all study areas

	Location	Latitude	Longitude		Location	Latitude	Longitude
	Location 1	29.625	-95.875		Location		
					1	19.125	-99.375
	Location 2	29.625	-95.625		Location		
					2	19.125	-99.125
	Location 3	29.625	-95.375	ation	Location		
					3	19.125	-98.875
	Location 4	29.625	-95.125		Location		
suo					4	19.375	-99.375
cati	Location 5	29.875	-95.875	Loc	Location		
Lo	Location 6	29.875	-95.625	ge]	5	19.375	-99.125
uge	Location 7	29.875	-95.375	gau	Location		
ga	Location 8	29.875	-95.125	PC	6	19.375	-98.875
cations and CPC	Location 9	30.125	-95.875	d CJ	Location		
	Location 10	30.125	-95.625	anc	7		
	Location 11	30.125	-95.375	ons		19.625	-99.375
	Location 12	30.125	-95.125	cati	Location		
				Lo	8	19.625	-99.125
ILC	Houston	29.75	-95.25	ΜI	Location		
N N	CPC gauge			RI	9		
IRI	Location 1			ty T		19.625	-98.875
on	Houston	29.75	-95.75	Ci	Mexico	19.25	-99.25
oust	CPC gauge			ico	City		
Hc	Location 2			Лех	CPC		
				4	gauge		
					Location		
					1		

					location	23.625	90.125
suo	location 1	-6.375	106.625	ions	1		
cati				cat	location	23.625	90.375
Loc	location 2	-6.375	106.875	, Lo	2		
Ige				nge	location	23.625	90.625
gau	location 3	-6.375	107.125	ga	3		
PC				JPC	location	23.875	90.125
qC	location 4	-6.125	106.625	D pu	4		
cations an				s ar	location	23.875	90.375
	location 5	-6.125	106.875	ion	5		
				cat	location	23.875	90.375
Lo	location 6	-6.125	107.125	ILc	6		
IM	Jakarta CPC	-6.25	106.75	MN	Dhaka	23.75	90.25
RN	gauge			IR	BMD		
ta T	Location 1			ka	gauge		
kar				Dha	Location		
Jaj					1		

Appendix VI: Standardized Precipitation Index Graphs

Standardized Precipitation Index Graphs for Houston







-1

-2

-3







 2006-

 2012-2013

E







Standardized Precipitation Index Graphs for Mexico City













Standardized Precipitation Index Graphs for Jakarta











Standardized Precipitation Index Graphs for Dhaka













Appendix VII: Urban Flood Risk Maps

[1314] Pinehurst Pinehurst ena Missouri City Missouri City August, 2007 August, 2017 Legend 30 I Kilometers 60 No Risk (≤ 0) Low Risk (0.01-6) Moderate Risk (6.01-12) High Risk (12.01-18) Extreme Risk (18.01-21)

Urban Flood Risk Map of Houston, Texas

Urban Flood Risk Map of Mexico City, Mexico





Urban Flood Risk Map of Jakarta, Indonesia

Tongi Tongi July, 2007 July, 2017 8.75 17.5 35 0 Legend Kilometers No Risk (≤ 0) Low Risk (0.01-6) Moderate Risk (6.01-12) High Risk (12.01-18) Extreme Risk (18.01-21)

Urban Flood Risk Map of Dhaka, Bangladesh

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